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Phytoplankton community in the Anthropocene: effects of climate change
and eutrophication

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Tese apresentada ao Programa de Pós-Graduação em Ecologia de Ambientes Aquáticos Continentais do Departamento de Biologia, Centro de Ciências Biológicas da Universidade Estadual de Maringá, como requisito parcial para obtenção do título de Doutor em Ecologia e Limnologia.

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*“E o fim é belo e incerto, depende de como
você vê...”*

Fernando Anitelle

Comunidade fitoplanctônica no Antropoceno: efeitos das mudanças climáticas e eutrofização

RESUMO

Na era do Antropoceno, os efeitos das mudanças climáticas nos ecossistemas aquáticos continentais apresentam múltiplos fatores, uma vez que o aquecimento e os diferentes fatores relacionados às mudanças climáticas influenciam os processos físicos, biogeoquímicos e biológicos. As mudanças climáticas causam impactos relevantes nos ecossistemas da Terra por meio do aumento da temperatura, mudanças nos padrões de precipitação, eventos climáticos extremos mais frequentes, mudanças nas concentrações de CO₂, entre outros. Os efeitos das mudanças climáticas somam-se aos efeitos contínuos de outros fatores globais e locais que afetam a diversidade, composição, estrutura e funcionamento das comunidades ecológicas, como eutrofização, fragmentação do *habitat*, mudanças no uso da terra e mudanças nos ciclos biogeoquímicos. Assim, esta tese é composta por dois segmentos. Para o primeiro segmento, desenvolveu-se um experimento para testar como o aumento das temperaturas influencia a diversidade fitoplanctônica e as emissões de CO₂ em ambientes eutróficos. Os resultados mostraram que, em cenários futuros de aquecimento, a composição da comunidade fitoplanctônica é alterada, afetando funções do ecossistema, como produção de biomassa, eficiência no uso de recursos e balanço de fluxo de carbono. O aquecimento agravou os efeitos negativos da eutrofização através do aumento das cianobactérias. Encontrou-se que a eutrofização pode promover mudanças climáticas, aumentando a liberação de gases de efeito estufa com evidências experimentais de um feedback positivo entre o principal sintoma de eutrofização (florações de cianobactérias) e o aquecimento, por meio de taxas de emissão de CO₂ mais altas em sistemas mais quentes dominados por cianobactérias, além de outras mudanças nas principais funções do ecossistema. No segundo segmento, conduziu-se um experimento de curto prazo para testar como diferentes comunidades fitoplanctônicas naturais, promovidas por diferentes regimes de temperatura, reagiram a um evento de chuva extrema simulada e, assim, analisar a estabilidade e resiliência do ecossistema. Constatou-se que as comunidades fitoplanctônicas submetidas a diferentes temperaturas responderam de forma diferente aos distúrbios. As comunidades dominadas por florações de cianobactérias se beneficiaram e mostraram-se resilientes ao evento de chuvas extremas. Em contraste, as comunidades supostamente menos estressadas pelo aquecimento tiveram uma resposta mais lenta ao evento de chuvas extremas e não recuperaram a biomassa antes do evento de chuvas extremas. Descobriu-se que as comunidades aparentemente mais estressadas (ou seja, sob a temperatura mais alta) têm maior estabilidade do ecossistema (resiliência, resistência e recuperação) quando comparadas às comunidades menos estressadas. Em suma, a evidência experimental indica que as mudanças climáticas afetarão profundamente a estrutura da comunidade e algumas funções do ecossistema (por exemplo, produção de biomassa, transferência de energia e ciclo do carbono). Destaca-se a necessidade de fortalecer as políticas e medidas locais para prevenir ou mitigar os impactos ecológicos das mudanças climáticas.

Palavras-chave: Mudanças climáticas. Múltiplos estressores. Eutrofização. Chuva. Cianobactéria.

Phytoplankton community in the Anthropocene: effects of climate change and eutrophication

ABSTRACT

In the era of the Anthropocene, the effects of climate change on the freshwater ecosystems are clearly complex, since warming and different temperature related drivers influence interacting physical, biogeochemical and biological processes. Climate change is predicted to have huge impacts on the Earth's ecosystems through temperature increase, changed patterns of precipitation, more frequent extreme weather events, and combinations of these thus, climate change may become one of the major drivers affecting the diversity, composition, structure, and functioning of ecological communities over the next several decades. In this thesis, composed of two papers, we evaluate the effects of multiple factors related to climate change on the natural phytoplankton community. In the first one, we conducted an indoor experiment to test how increasing temperatures influence natural phytoplankton diversity and CO₂ emissions in eutrophic ecosystems. Our results experimentally show that, under future scenarios of climate warming, the phytoplankton community composition can respond strongly, affecting ecosystem functions such as biomass production, resource use efficiency, carbon flux balance. Warming clearly aggravated the negative effects of eutrophication through the enhancement of cyanobacteria, all other factors being equal. Since the suggestion that eutrophication may promote climate change by increasing the release of greenhouse gases from fresh waters, it has been found that eutrophication may interact with warming via a positive feedback to atmospheric CH₄ emissions. Here, we also found experimental evidence of a positive feedback between the major eutrophication symptom (cyanobacterial blooms) and warming, via higher CO₂ emission rates in cyanobacteria dominated warmer systems, besides other changes in key ecosystem functions. In the second paper, we conducted an indoor short-term experiment to test how the natural phytoplankton community subjected to different temperatures reacted to the stressors of climate change (warming, eutrophication, extremes rainfall events) can affect the ecosystem stability. We find that the phytoplankton communities responds differently to disturbances. The environments with cyanobacterial blooms have benefited and proved to be resilient to the extremes rainfall events. In contrast, environments less stressed by warming have a slower response to the event of extreme rainfall, and that they often do not recover their biomass before the extreme rainfall event. Thus, given the multiple effects of climate change, the most stressed environments have greater ecosystem stability (resilience, resistance, and recovery) when compared to the least stressed. In summary, the findings of this thesis, we have experimental evidence with phytoplankton community, that climate change will profoundly affect ecosystem functions (e.g., biomass production, energy transfer, and carbon cycle). In this sense, we hope to contribute with policies to prevent or mitigate the ecological impacts of climate change.

Keywords: Climate change. Multiple stressors. Eutrophication. Rainfall. Cyanobacteria.

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1 **1 GENERAL INTRODUCTION**

2
3 The Anthropocene term informally encompass different geological, ecological,
4 sociological, and anthropological changes in recent Earth history (Malm and Hornborg, 2014;
5 Steffen et al., 2011). Climate change is predicted to have huge impacts on the Earth's
6 ecosystems through temperature increase, changed patterns of precipitation, more frequent
7 extreme weather events, and combinations of these (Field *et al.* 2014, Lehmann *et al.* 2015).
8 Thus, climate change may become one of the major drivers affecting the diversity,
9 composition, structure, and functioning of ecological communities over the next several
10 decades.

11 Lakes are considered good sentinels of climate change because they are sensitive to
12 environmental changes and can integrate changes in the surrounding landscape and the
13 atmosphere (Adrian et al., 2009). Besides climate change, freshwater ecosystems are
14 threatened by multiple anthropogenic stressors that include, for example, habitat
15 fragmentation and isolation, overexploitation, invasion by exotics, and eutrophication (Yvon-
16 Durocher et al., 2011; Meerhoff et al., 2012; Birk et al., 2020; Albert et al., 2021).
17 Biodiversity losses (e.g. due to eutrophication) is occurring as a result of global changes
18 putting the functioning of aquatic ecosystems at risk (Cardinale et al., 2006; Dudgeon et al.,
19 2006; Hooper et al., 2012). Together with global warming and rising atmospheric CO₂ levels,
20 these pressures are altering life on Earth in unpredictable ways, with potentially very severe
21 consequences for the goods and services that ecosystems provide for humanity (Steffen et al.,
22 2015). Therefore, understanding the ecosystem responses and stability to these stressors is
23 crucial to be able to provide better management, conservation and restoration strategies
24 (Scheffer et al., 2001; Pecl et al., 2017; De Boeck et al., 2018).

25 Phytoplankton community is extremely sensitive to environmental change, responding
26 with changes in total biomass and community composition (Litchman et al., 2015). Both

27 changes in temperature and precipitation patterns can strongly affect phytoplankton in direct
28 and indirect ways. It has been predicted that warming will increase the occurrence of blooms
29 (Paerl and Huisman, 2009, 2008; Kosten et al., 2012, Medeiros et al., 2015), or at least favor
30 cyanobacterial dominance within phytoplankton communities (Gkelis et al., 2014; Yan et al.,
31 2017). Climate change is also increasing the variability and extremeness of precipitation and
32 its impacts on blooms are not well understood. Extreme rainfall events change lake abiotic
33 conditions. The physical displacement of phytoplankton throughout the water column can
34 alter the outcome of competition and herbivory and thus shape community composition
35 (Reynolds et al., 2002; Paerl & Huisman, 2009).

36 The cyanobacterial blooms are favored because they exhibit a series of adaptations to
37 enable survival in a range of extreme niches (Winder & Sommer, 2012) and are generally
38 difficult for higher trophic-level consumers to assimilate. Other factors that allow the success
39 of cyanobacteria are functional characteristics that increase their fitness in a wide range of
40 environmental characteristics (Litchman et al., 2010). These characteristics include the
41 presence of aerotopes that allow them to move through the water column when there is
42 thermal stratification, the nitrogen-fixing capability, high affinity and capacity of stocking
43 phosphorous, and the production of cysts (akinetes) (Hansson, 2000; Weyhenmeyer et al.,
44 2007; Carey et al., 2012). Thus, it is expected that cyanobacteria blooms become more
45 frequent and widespread in the future given the expansion of eutrophication worldwide and
46 climate warming, as indicated by theoretical models (Mooij et al., 2005; O'Neil et al., 2012)
47 and empirical data (De Senerpont Domis et al., 2007; Kosten et al., 2012). An expansion of
48 cyanobacterial blooms is of great societal concern because harmful cyanobacteria can impair
49 safe drinking, irrigation, fishing and recreational waters that are critical for the growing global
50 human population (Heino et al., 2020).

51 In recent years, numerous studies have indicated that eutrophication, rising CO₂ levels,
52 and global warming are likely to interact additively or synergistically to increase the
53 frequency, intensity, and duration of cyanobacterial blooms in many aquatic ecosystems
54 across the globe (Paerl & Huisman, 2009; Wagner & Adrian, 2009; Qin et al., 2015; Yan et
55 al., 2017). Although nutrients seem to be the most important predictor of cyanobacterial
56 biovolume, as lakes become more eutrophic cyanobacteria become more sensitive to the
57 interaction of nutrients and temperature (Rigosi et al., 2014).

58 Theoretical and experimental evidence highlight the potential reinforcing feedbacks
59 between eutrophication and warming (Davidson et al., 2018; Moss et al., 2011; Yan et al.,
60 2017; Li et al., 2021), among other mechanisms by altering the ‘metabolic balance’ of
61 ecosystems (Allen et al., 2005) and impacts on the fluxes of greenhouse gases such as CH₄
62 and N₂O, besides CO₂. The metabolic balance is defined as the rate between carbon fixed
63 through photosynthesis and its remineralization through respiration, determining whether an
64 ecosystem acts as a net source or sink of CO₂ to the atmosphere (Del Giorgio & Duarte, 2002;
65 Odum, 1956; Woodward, 2007). In this sense, some studies show that respiration responds
66 more strongly to temperature change than photosynthesis (Gudasz et al., 2010; Moss et al.,
67 2010; Yvon-Durocher et al., 2011). Current studies indicate that, in contrast to previous
68 beliefs, freshwater ecosystems are more active in terms of carbon sequestration, processing,
69 and burial than terrestrial and marine ecosystems (Downing, 2010; Tranvik et al., 2009;
70 Raymond et al., 2013). Climate warming has been shown to increase greenhouse emissions
71 and reduce carbon sequestration in these environments (López-Urrutia et al., 2006; Yvon-
72 Durocher et al., 2011) (López-Urrutia et al., 2006; Yvon-Durocher et al., 2011). These
73 findings are important because they imply carbon cycle responses to climate warming should
74 be more complex than the simple temperature effect on respiration rates alone. Besides carbon
75 cycle, climate warming can affect other ecosystem processes. Among them, resource use

76 efficiency (RUE) is a key indicator of ecosystem functioning, since it resumes nutrient
77 cycling and trophic transfer processes (Ptacnik et al., 2008; Filstrup et al., 2014).

78 The role of experiments, and of microcosms in particular, might be crucial to explore
79 the effects and the consequences of multiple stressors, such as climate change and the effects
80 of biodiversity loss on ecosystem functioning. Although the use of model systems might
81 appear to be limited in scope and realism, especially compared with the infinite complexities
82 of the real world and the spatial extent of global-scale problems, the utility of the microcosm
83 approach lies in its ability to explore and test mechanisms (Benton et al., 2007). Together with
84 other sources of scientific knowledge, microcosm experiments can supply robust scientific
85 evidence to base measures and policies to prevent or mitigate the ecological impacts of
86 environmental change.

87 In this sense, the experimental studies approach can offer a tool to reach “The
88 Sustainable Development Goals”, proposed by the United Nations (UN) and adopted by all
89 UN Member States in 2015, whose action is to protect the planet and improve the lives and
90 prospects of everyone, everywhere. The 17 Goals is part of the 2030 Agenda for Sustainable
91 Development. The development of this thesis directly contemplates the dimension of the
92 Biosphere but with economic and social projections, such as Goal 6: Ensure access to water
93 for all, and Goal 13: Take urgent action to combat climate change and its impacts. The
94 complexity of issues about climate change affects several individuals and social groups,
95 requiring the participation of a greater number of social actors in the search of possibilities
96 and alternatives for adaptation and mitigation for the imposed uncertainties. Thus, scientific
97 dissemination is an indispensable tool for the connection between researchers and society,
98 because the individual's understanding of the world directly influences their choices and
99 decisions (Supporting Information Fig.1).

100 The general aim of this thesis was to test how natural freshwater phytoplankton
 101 diversity responds to components of climate change and how this response may influence
 102 some ecosystem processes. To address this general objective, two in-door microcosm
 103 experiments were conducted. Specifically, in the first one, we tested how increasing
 104 temperatures influenced natural phytoplankton diversity and CO₂ emissions in eutrophic
 105 conditions. The potential links and feedback between eutrophication and warming were thus
 106 explored. In the second paper, we conducted an indoor short-term experiment to test how
 107 different natural phytoplankton communities, acclimatized to different temperatures, reacted
 108 to a simulated extreme rainfall event, and thus analyze ecosystem stability and resilience. We
 109 expect that under higher temperatures, rainfall extreme events in eutrophic lakes will result in
 110 increased cyanobacteria blooming.

111

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- 256

257 2 POSITIVE FEEDBACK BETWEEN WARMING AND CYANOBACTERIA 258 BLOOMS

259

260

261 ABSTRACT

262 The synergistic effect of global warming and eutrophication favors the formation of
263 blooms of cyanobacteria and alters the functioning of ecosystems (e.g. biogeochemical
264 cycles and productivity). In this study, we evaluated the potential feedbacks between
265 eutrophication and warming on the phytoplankton and how global warming will affect
266 the metabolic balance. We conducted an indoor controlled microcosm experiment.
267 Warming promoted species richness decreased and primary production with a 10-fold
268 increase in the mean biomass of green algae and cyanobacteria (*Raphidiopsis*
269 *raciborskii*). Resource use efficiency (RUE) increased gradually between treatments
270 during the experimental .Maximum RUE value was obtained under the warmest
271 treatment, dominated by few cyanobacteria genera which are more efficient in resource
272 use and limiting the growth of other species. Although we registered high CO₂ influx
273 values in the experiment, the ecosystem metabolic balance changed with temperature
274 increase. The magnitude of influx decreased with warming, almost transforming the
275 microcosms in sources of CO₂ to the atmosphere. We designed our experiment to focus
276 on increasing mean temperature as a driver of changes in the phytoplankton community,
277 providing non-limiting nutrients and no predation conditions, to avoid indirect or
278 interacting effects of warming. Therefore, we can assume that warming directly drove
279 cyanobacteria proliferation. Here, we also found experimental evidence of positive
280 feedback between eutrophication symptoms (blooms) and warming, via higher CO₂
281 emission rates in cyanobacteria dominated warmer systems. Thus, we were able to
282 capture a pure pelagic response to warming, exclusive to the phytoplankton, and how

283 such responses translated to ecosystem functions (e.g., biomass production, energy
284 transfer, and carbon cycle).

285

286 **Keywords:** Climate change, cyanobacteria blooms, eutrophication, global warming,
287 resource use efficiency, feedback.

288

289 2.1 Introduction

290 Eutrophication of surface waters has become a ubiquitous problem around the
291 world, threatening both aquatic biodiversity and several services for the human
292 population, such as water supply, recreation, and irrigation (Heino et al., 2020; Jeppesen
293 et al., 2009; Moss et al., 2009). Climate change may enhance the negative effects of
294 eutrophication through a variety of physical, chemical, and biological mechanisms
295 (Moss et al., 2011). Warming affects biogeochemical cycles and biological processes
296 such as respiration and decomposition rates, nutrient cycling, growth rates, individual
297 and community body size, and environmental selection of functional traits, among other
298 ecological processes (Jeppesen et al., 2010; Meerhoff et al., 2012; Gkelis et al., 2014).
299 As a result, there is selective pressure on aquatic organisms, such as on groups of
300 phytoplankton that can withstand such changes (Mouillot et al., 2013b).

301 Across most of the studies from laboratory and field observations, there was a
302 general trend of enhanced cyanobacteria biomass and/or dominance with increasing
303 water temperature. Rasconi *et al.* (2017) designed a mesocosm experiment and found a
304 clear effect of the temperature treatments with an observed shift toward cyanobacteria
305 dominance. However, some empirical studies have shown that not the temperature *per*
306 *se*, but the interaction with the nutrient supply is important for increasing the biomass of
307 these organisms (Thrane et al., 2017; Verbeek et al., 2018). There are still contrasting
308 experimental evidence that showed no particular pattern of the cyanobacteria with rising
309 temperature, at least in shallow unstratified lakes still dominated by macrophytes (Moss
310 et al., 2003) or that the growth of cyanobacteria in a global warming scenario does not
311 exceed the growth rates of green algae (Lürding et al., 2013). Given the wide spectrum
312 of climate change scenarios (Moss et al., 2010), predicting global warming effects on
313 aquatic ecosystems is still a challenge with a high level of uncertainty (Feuchtmayr et

314 al., 2009). Despite this, the increase in the magnitude and frequency of cyanobacteria
315 blooms associated with higher temperatures is a global concern (Paerl & Huisman,
316 2008; Kosten et al., 2012; Kruk et al., 2012; Paerl & Paul, 2012; Michalak, 2016;
317 Burford et al., 2019).

318 The competitive advantage that allows the success of some cyanobacteria groups
319 are functional characteristics that increase their fitness over a wide range of
320 environmental characteristics (Litchman et al., 2015). These characteristics include the
321 presence of aerotopes that allow them to move through the water column when there is
322 thermal stratification, nitrogen-fixing capacity, high affinity for and capacity of stocking
323 phosphorous, and the production of cysts (akinetes) (Carey, Ibelings, Hoffmann,
324 Hamilton, & Brookes, 2012; Hansson, 2000; Litchman, Pinto, Klausmeier, Thomas, &
325 Yoshiyama, 2010; Weyhenmeyer et al., 2007). Thus, it is expected that cyanobacterial
326 blooms become more frequent in a global warming scenario (Mooij et al., 2005; O'Neil
327 et al., 2012; Beaulieu et al., 2013). Biodiversity losses (e.g. eutrophication) is occurring
328 as a result of global climate change and puts at risk the functioning of the aquatic
329 ecosystems (Cardinale et al., 2006; Dudgeon et al., 2006; Hooper et al., 2012). Among
330 ecosystem functions, resource use efficiency (RUE) is a very important one,
331 determining nutrient cycling, and trophic transfer processes (Ptacnik et al., 2008;
332 Filstrup et al., 2014).

333 Theoretical and experimental evidence highlight the potential feedbacks between
334 eutrophication and warming (Davidson & Janssens, 2006; Moss et al., 2011; Yan et al.,
335 2017), among other mechanisms by altering the 'metabolic balance' of ecosystems
336 (Allen et al., 2005). This balance is defined as the rate between carbon fixed through
337 photosynthesis and its remineralization through respiration, determining whether an

338 ecosystem acts as a net source or sink of CO₂ to the atmosphere (Odum, 1956; Del
339 Giorgio & Duarte, 2002; Woodward, 2007).

340 Eutrophication may decrease the relative importance of the external organic
341 matter and promote higher autotrophic fixation of CO₂ by cyanobacterial blooms,
342 transforming the lakes into net carbon sinks. The synergistic effect of warming and
343 eutrophication can also promote higher respiration of settling the organic matter,
344 promoting the release of CO₂ from the lakes (Gudasz et al., 2010; Moss et al., 2010,
345 2011; Yvon-Durocher et al., 2010, 2011). Therefore, ecosystem respiration is affected by
346 temperature more than photosynthesis rates and, without other interacting factors,
347 warming may increase CO₂ emissions and reduce net carbon sequestration by eutrophic
348 aquatic ecosystems (López-Urrutia et al., 2006; Yvon-Durocher et al., 2010). Still, it
349 remains uncertain to what extent changes in community composition as a response to
350 warming directly translate into changes in ecosystem function, such as CO₂ fluxes.

351 Here, we conducted an indoor experiment to test how increasing temperatures
352 influence natural phytoplankton diversity and CO₂ emissions in eutrophic ecosystems.
353 We expected that, as systems were isolated without immigration, water temperature
354 increase would reduce phytoplankton richness and promote changes in composition,
355 leading to a rapid dominance of some phytoplankton groups, mainly cyanobacteria.
356 Furthermore, as nutrients were not limiting, we expected that warmer waters would
357 promote higher total biomass and different functional responses, such as higher
358 respiration rates increasing net CO₂ effluxes, and higher RUE. In summary, we expect
359 that under higher temperatures eutrophic lakes will become dominated by cyanobacteria
360 and act as net carbon sources to the atmosphere, creating a positive feedback between
361 eutrophication and warming.

362 2.2 Methods

363 2.2.1 Experimental Design

364 We conducted an indoor controlled microcosm experiment in a laboratory
365 located at the *campus* of the Universidad de la Republica (UdelaR), Maldonado,
366 Uruguay (34° 54'53'S e 54 ° 56'31 " W '), between June 18 and July 19 of 2019
367 (southern hemisphere winter). The control treatment, (i) 17 °C, corresponded to the
368 annual mean temperature of the aquatic ecosystems in the same region where the
369 experiment was conducted (Pacheco et al., 2010). The second treatment, (ii) 20 °C,
370 represented a scenario with an increase of 3 °C in relation to the annual mean
371 temperature; and the third treatment, (iii) 23 °C, represented an increase of 6 °C in
372 relation to the mean temperature. Room temperature was manipulated in all treatments
373 using electronic temperature controllers (Eliwell ID Plus 961), including the control
374 treatment.

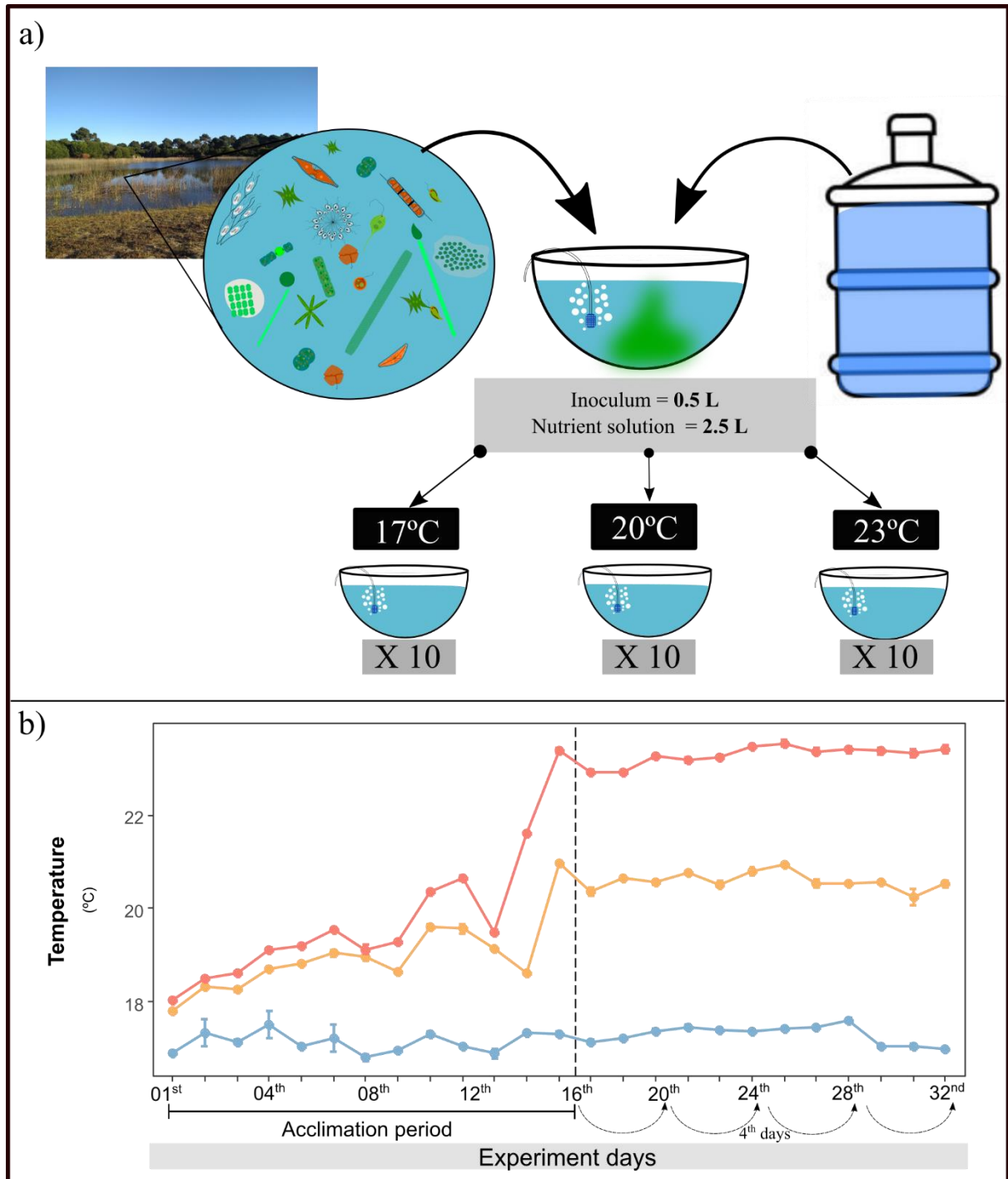
375 The microcosm was cylindrical polyethylene 5-L aquaria (23.5 cm of diameter
376 and 10 cm of height). We installed water circulation pumps (model Resun AC-9903) at
377 the bottom of each aquarium to avoid phytoplankton sedimentation (Flury et al., 2010;
378 Sommer et al., 2015). Besides, we moved the water manually and delicately twice a day
379 with a stick. Both light and temperature were controlled. Phytoplankton received circa
380 80 lum/ft² of 12/12 ratio of light/darkness from fluorescent lamps with a light spectrum
381 similar to that of the sun. All experimental units were placed at the same height to avoid
382 temperature fluctuations and to homogenize the quantity of incident light. We randomly
383 distributed thirty microcosms between the three temperature treatments (i.e. n=10
384 replicates each).

385 To prepare the natural phytoplankton inoculum for the experiment, we obtained
386 water samples at the subsurface of the limnetic region in a series of subtropical shallow
387 lakes situated at the Uruguayan coastline (34° S 56° W). These lakes comprise a trophic

388 gradient from mesotrophy to hypertrophy (Kruk et al., 2009; Pacheco et al., 2010). We
389 collected phytoplankton using a plankton net with a mesh size of 20 μm to remove most
390 zooplankton and added some non-filtered water to include smaller phytoplankton
391 species ($< 20 \mu\text{m}$). We used this strategy to maximize the sampling of less abundant
392 species and guarantee that most taxonomic groups were present in the experimental
393 units. The mean water temperature measured in the lakes during the sampling procedure
394 was 15 $^{\circ}\text{C}$.

395 We distributed 0.5 L of concentrated inoculum for each of the 30 experimental
396 units and added 2.5 L of dechlorinated water enriched with a nutrient solution to reach a
397 final volume of 3 L. In the first day of the experiment, we found no significant
398 differences among the three treatments for phytoplankton taxonomic richness (F-value
399 $(2, 27) = 1.031$; $p = 0.37$) and biomass (F-value $(2, 27) = 0.086$; $p = 0.918$), showing the
400 homogeneous phytoplankton distribution among the treatments.

401 To avoid losses due to heat shock due to abrupt temperature changes, the
402 experimental units were acclimated for two weeks after the addition of the
403 phytoplankton by increasing room temperature gradually and daily until final
404 temperatures were reached (approximately 0.5 $^{\circ}\text{C}$ increase per day), as this allowed for
405 changes in species abundance in response to the temperature increase but avoided heat
406 shocks. Thus, on the 16th day of the experiment, all climatic rooms had achieved the
407 final temperatures for each treatment. The experiment lasted for 32 days, as the short
408 life cycle of phytoplankton allows the reproduction of several generations in a few days
409 (Reynolds, 2006) (Fig 1, Supplementary material S1).



410

411 **Figure 1:** Experimental design. a) A concentrated phytoplankton community
 412 subtropical lakes were cultured in a nutrient solution respecting the Redfield ratio.
 413 Treatments: 17 °C, 20 °C and 23 °C, total n=30. b) Following an acclimation period, as
 414 from the 16th day of the experiment all treatments reached the wished temperature. The
 415 central point denotes the mean value and, whiskers represent standard error for each
 416 temperature in each experiment day.

417 Phytoplankton was maintained in a medium following the Redfield ratio. We
418 prepared a solution composed of 200 mg L⁻¹ KH₂PO₄, 100 mg L⁻¹ NH₄NO₃, 177 mg L⁻¹
419 Ca(NO₃)₂, 0.1 mg L⁻¹ Co(NO₃)₂, and 250 mg L⁻¹ C₁₀H₁₆N₂O₈. To this solution, we added
420 150 L of deionized water and, after intense homogenization, added 2.5 L of it to each
421 microcosm. Weekly, the same solution was added to each microcosm to compensate for
422 the losses due to evaporation and to reach the initial volume of 3 L (McKee et al., 2000;
423 Ekvall & Hansson, 2012). On average, each microcosm presented 533.23 µg L⁻¹ of TN
424 and 81.2 µg L⁻¹ of TP (Supplementary material Fig.S1).

425 2.2.2 Limnological variables

426

427 In all microcosms, we took daily measurements of water temperature (°C), pH,
428 and electric conductivity using a HANNA multiparametric probe, and of dissolved
429 oxygen (mg L⁻¹) and oxygen saturation (%) using an Oxyguard Handy Polaris. We
430 sampled water on the 4th, 16th, 24th, and 32nd day to determine total phosphorous (TP; µg
431 L⁻¹), total nitrogen (TN; µg L⁻¹), reactive soluble phosphorous (SRP; µg L⁻¹), nitrate
432 (NO₃; µg L⁻¹), and ammonium (NH₄⁺; µg L⁻¹), as well as chlorophyll-a, according to
433 (APHA, 2005). The limnological variables TP, TN, SRP, NO₃, NH₄⁺ were log-
434 transformed. We found no significant differences in nutrient concentration for our
435 treatments through time, guaranteeing that these conditions were successfully controlled
436 throughout the experiment (Supplementary material Fig.S2).

437

438 2.2.3 Phytoplankton

439

440 We sampled phytoplankton every four days (days 1, 4, 8, 12, 16, 20, 24, 28, and
441 32). Before taking each sample, we homogenized the water manually to avoid missing
442 any species due to sedimentation. We sampled phytoplankton directly with flasks and

443 fixed them immediately with acetic Lugol solution. We estimated phytoplankton density
444 following Utermöhl (1958) and Lund et al., (1958) and calculated density according to
445 APHA (2005). The biomass ($\text{mm}^3 \cdot \text{L}^{-1}$) was considered as biovolume, which was
446 estimated by multiplying the density of each taxon by its volume. We estimated the cell
447 volume by calculating the volume of the geometric shape that was the most similar to
448 each cell form (Sun & Liu, 2003). We also estimated the community resource use
449 efficiency (thereafter RUE), defined as the ratio between the phytoplankton biomass
450 production in Chl-a and TP, as a proxy for ecosystem productivity (Ptacnik et al., 2008;
451 Olli et al., 2015; Verbeek et al., 2018).

452

453 2.2.4 Data analyses

454

455 We used 1-way ANOVA to test for differences among the treatments in nutrient
456 concentration (TN, TP, SRP, NO_3 , and NH_4) in each sampling occasion; in
457 phytoplankton richness and biomass on the first day of the experiment; and two-way
458 ANOVAs to test for differences in chlorophyll-a, RUE, and carbon fluxes among
459 treatments and time. When results were significant ($p < 0.05$), we used Tukey tests, to
460 verify the significant difference in mean among treatments and time. After the
461 temperature was stabilized in all treatments (as from the 16th day), we evaluated the
462 effects of the temperatures and time on richness and biomass of phytoplankton groups
463 by fitting generalized linear models (GLM), using the distribution families that better
464 adjusted to the data: Poisson distribution for richness and Gaussian distribution for
465 biomass. We evaluated overdispersion and corrected them when necessary.

466 To evaluate the effects of temperature (three levels) and sampling time (five
467 levels) on the composition of phytoplankton (presence/absence) (after the 16th day), we
468 used a permutational multivariate analysis of variance (PERMANOVA, Anderson,

2001). The test was based on a Jaccard dissimilarity matrix, and 999 permutations were used to test for significance. To verify patterns in species composition through time we represented our data using an analysis of principal coordinates (PCoA)(Legendre & Legendre, 1998).

The rate of net growth (NG) of the phytoplankton density was calculated per phytoplankton group as density: $\mu = (\ln N_F - \ln N_0)/t$, where t is the time of the experiment (32 days), $\ln N_F$ is the natural logarithm of the density of organisms at the end of the experiment (32nd day), and $\ln N_0$ is the natural logarithm of the density of organisms at the beginning of the experiment (1st day).

The CO₂ fluxes were measured on the 10th, 16th, 24th, and 32nd day after the beginning of the experiment in all microcosms ($n = 120$), using an adapted environmental CO₂ analyzer (EGM-4) with a hermetic acrylic cover. We turned the aerators off one hour before we began sampling with the EGM-4, to avoid intervention in the CO₂ emission. We sampled data in each microcosm at a frequency of 30 seconds for 5 minutes, so that we obtained 10 measurements of [CO₂] per microcosm in the light period (photosynthesis period) and 10 in the dark period (respiration period). We calculated the balance of each day of CO₂ evaluation with the equation:

CO₂ emission of the sampled time = Photosynthesis + Respiration.

We used the slope of the relationship between gas concentration and time to calculate the gas flux according to:

$$\text{CO}_2 \text{ flux} = S * \frac{V}{A} * (P * F1 * F2 / R * T)$$

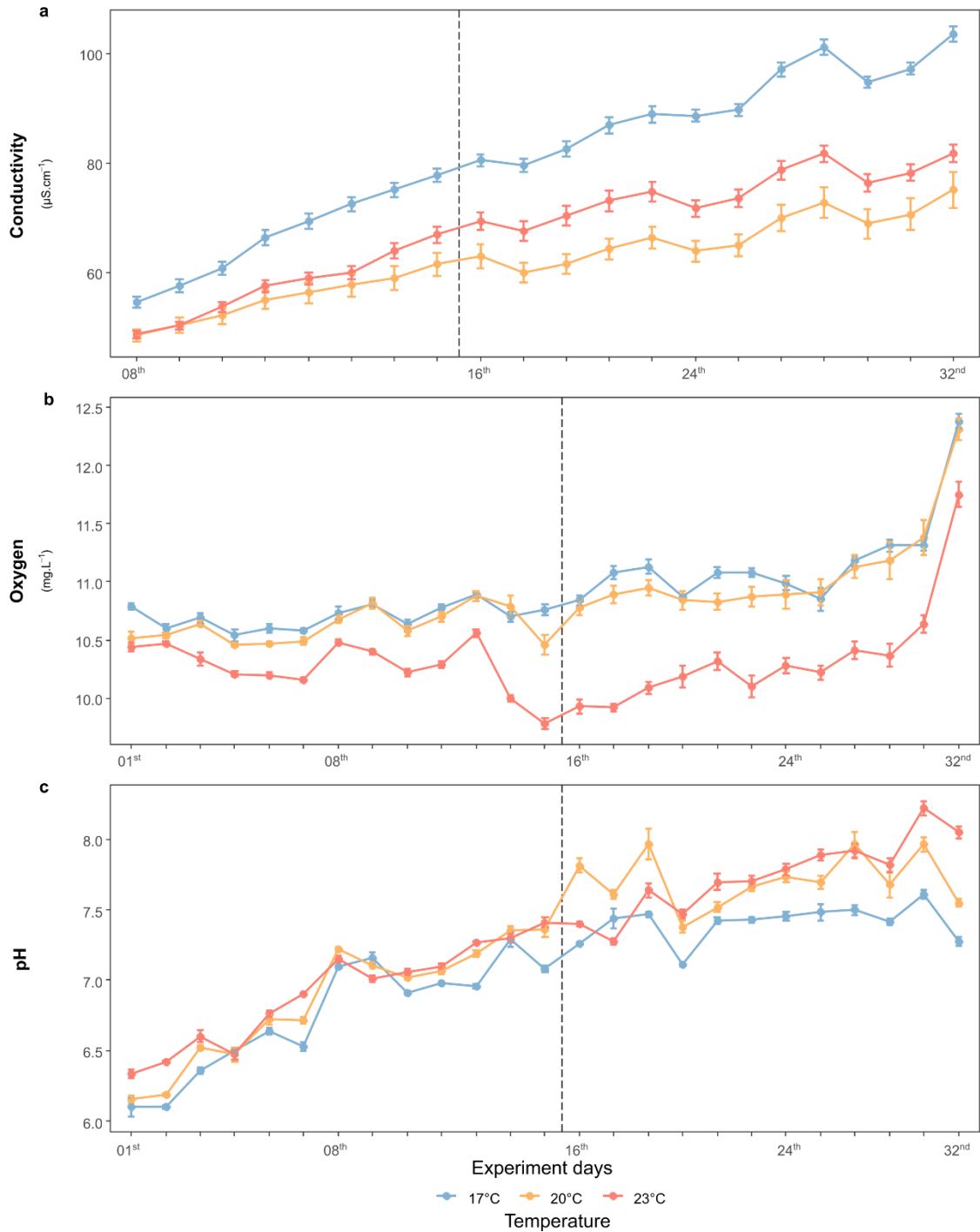
Where CO₂ flux is in CO₂.mg⁻².d⁻¹, S is the slope of the partial pressure of CO₂ in the function of time (ppm.s⁻¹), V is the camera volume (0.0019 m³), A is the camera area (0.1589 m²), P is the atmospheric pressure in KPa (102.5), $F1$ is the molecular weight of the gas (44g.mol⁻¹), $F2$ is the factor of conversion of seconds to days (8,64.104 s.d⁻¹), R is the gas constant

494 $(8.31\text{J}\cdot\text{mol}^{-1}\cdot\text{K}^{-1})$ and T is the air temperature in Kelvin. We selected values of the slope only
495 for those relationships with $R^2 \geq 0.7$. Positive values indicate CO_2 efflux (CO_2 liberation to
496 the atmosphere: predominance of net respiration) and negative values indicate CO_2 influx
497 (CO_2 retention in the aquatic environment: predominance of net photosynthesis). We ran all
498 analyses in software R version 3.3.2 (R Development Core Team, 2021), using packages
499 “vegan” (Oksanen et al., 2016) for PCoA, ANOVA, and PERMANOVA, and “MASS”
500 (Venables & Ripley, 2002) for GLMs.

501 **2.3 Results**

502

503 To electric conductivity, there was a gradual increase through time in all
504 treatments. The highest mean values were recorded for the treatment of 17°C ($81\ \mu\text{S}\cdot\text{cm}^{-1}$)
505 and minors for the treatment of 20°C ($62.1\ \mu\text{S}\cdot\text{cm}^{-1}$). The concentrations of dissolved
506 oxygen and pH values had little variation over the time of the experiment. However, the
507 warming promoted environments with lower average values of oxygen concentrations
508 ($10.31\ \text{mg}\ \text{L}^{-1}$) and higher mean pH values (8.05) in relation to the lower temperature
509 treatment ($10.95\ \text{mg}\ \text{L}^{-1}$ e 6.1 respectively).



510

511 **Figure 2:** Variation of main limnological variables in the three treatments (17 °C, 20
 512 °C and 23 °C) through the 32 days of the experiment. The dotted line indicates that on
 513 the 16th day of the experiment the wished temperatures were reached. The central point
 514 denotes the mean value and, whiskers represent standard error for each temperature in
 515 each experiment day.

516

517 2.3.1 Phytoplankton community

518

519 We identified 125 taxa, including green algae (62 taxa), cyanobacteria (21),
520 diatoms (18), and phytoflagellates (16). *Scenedesmus* (Chlorophyceae) presented the
521 highest number of taxa (7). Mean phytoplankton richness presented little oscillation
522 through the duration of the experiment (Fig.S2). However, we registered a reduction in
523 species richness with increasing temperature for total phytoplankton (pseudo $R^2 = 0.16$;
524 $p = 0.001$) and also and within the main phytoplankton groups (Table 1, Supplementary
525 material Fig. S3). The highest variability in species composition occurred in the 20 °C
526 treatment and the lowest in the 17 °C treatment (Fig. 3). The separation detected in the
527 PCoA was confirmed by the PERMANOVA test, which showed significant changes in
528 phytoplankton composition with temperature ($R^2 = 0.54$; $p = 0.001$) and time ($R^2 =$
529 0.07 ; $p = 0.001$).

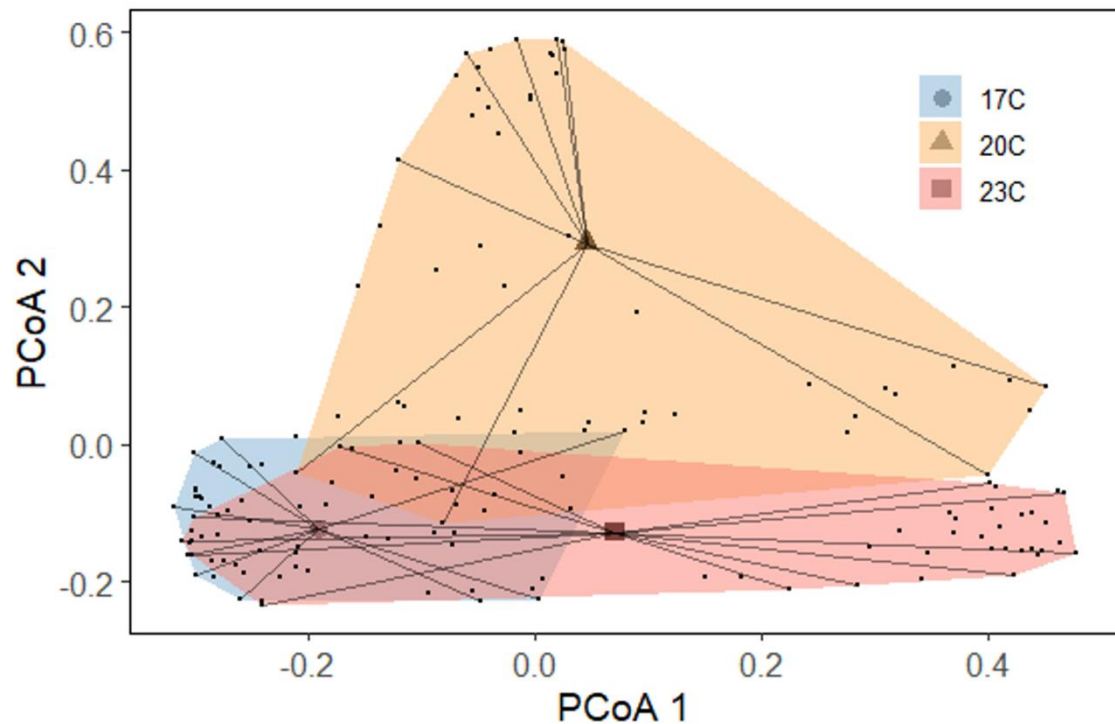
530 Cyanobacteria (as *Raphidiopsis raciborskii* (Wołoszyńska) Aguilera, Berrendero
531 Gómez, Kastovsky, Echenique & Salerno, before *Cylindrospermopsis raciborskii*), and
532 green algae (as *Desmodesmus magnus* (Meyen) Tsarenko and *Staurastrum* sp.) were the
533 groups that contributed to phytoplankton biomass the most through the experiment in
534 the three treatments compared to a more balanced community at the start of the
535 experiment. Supporting our expectations, total phytoplankton biomass responded
536 positively to warming (Pseudo $R^2 = 0.50$; $p < 0.001$) (Table 2, Fig.4, Supplementary
537 material Fig.S4).

538 This result was mainly driven by higher cyanobacteria growth at higher
539 temperatures (Pseudo $R^2 = 0.41$; $p < 0.001$). Considering the three treatments evaluated,
540 between the first and the last experiment day, we registered a 10-fold increase in the
541 mean biomass of green algae and cyanobacteria and the highest growth rates in the 20°
542 and 23 °C treatments (Fig. 5). Thus, warming promoted increasing of the phytoplankton

543 biomass (F-value $(2, 85) = 25.95$; $P < 0.001$) (Fig. 6a) reflected by higher mean values of
 544 chlorophyll-a concentrations in the warmer treatments of the experiment over time (F-
 545 value $(2, 85) = 12.23$; $P < 0.001$).

546 The RUE changed significantly over time and between the treatments during the
 547 experiment. RUE increased gradually between treatments during the experimental
 548 period and the maximum was obtained under the warmer treatment (F-value $(2, 57) =$
 549 13.40 ; $P < 0.001$). Furthermore, the blooms registered in this treatment, mostly of *R.*
 550 *raciborskii* at 23°C showed higher RUE than communities in the low temperature
 551 treatment. The differences being the largest between 17°C and 23°C (Tukey test $p =$
 552 0.04) (Fig. 6b).

553



554

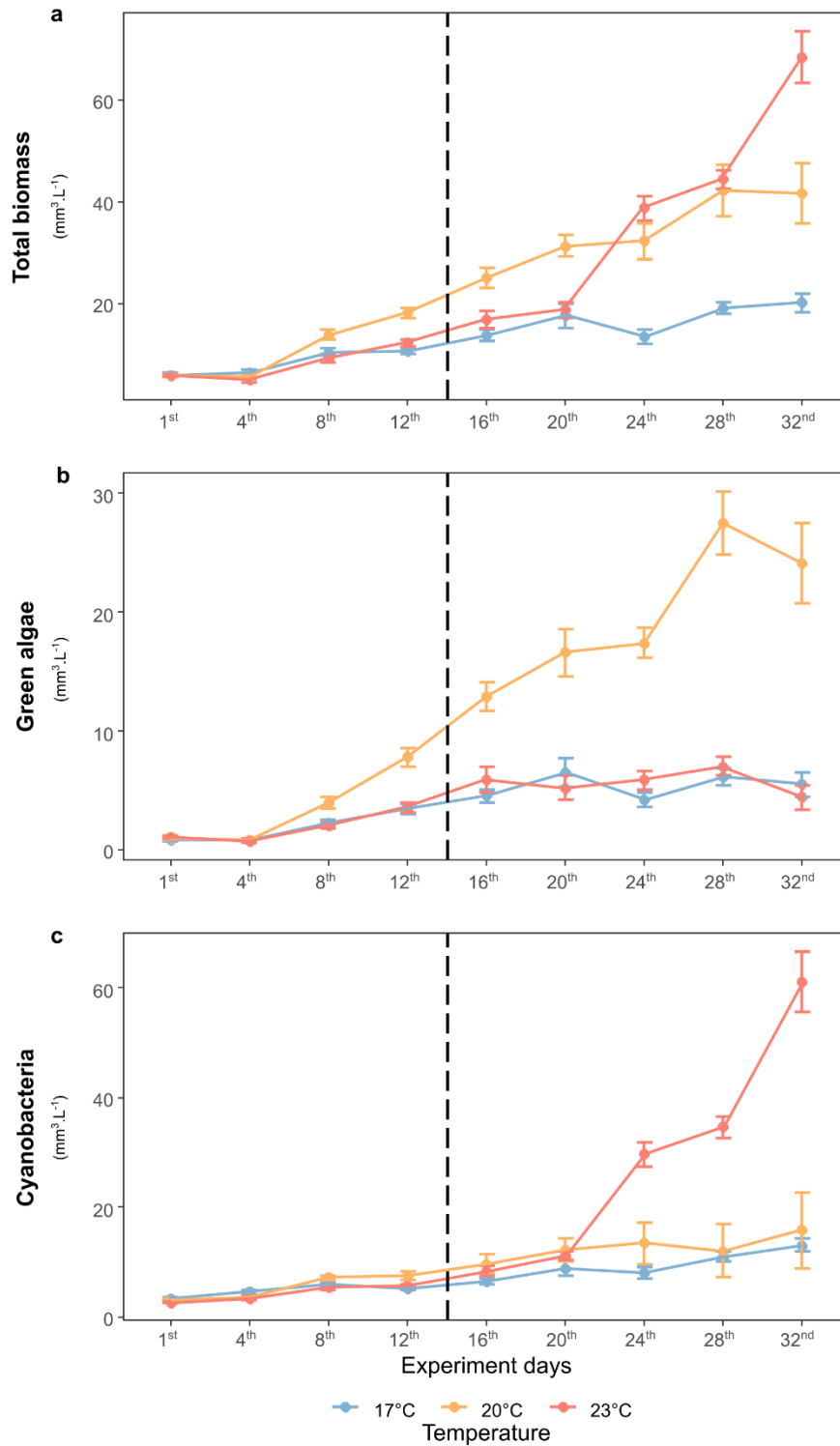
555 **Figure 3.** Principal coordinate analysis (PCoA), showing the variability in the
 556 composition of phytoplankton species from the 16th day of the experiment in the three
 557 treatments.

558

559 **Table 1** Generalized linear models, indicating regression estimate, standard errors (SE),
 560 z-value, and P-values of models predicting, as from the 16th day, the taxonomic
 561 richness of total phytoplankton, green algae, and cyanobacteria. Significant results are
 562 shown in bold.

Response variables	Predictors variables	Estimate	SE	z-value	<i>P</i>	Pseudo R ²
Total Richness	Intercept	3.907	0.164	23.72	0.000	0.16
	Temperature	-0.033	0.007	-4.537	0.000	
	Time	-0.008	0.003	-2.440	0.015	
Green algae	Intercept	3.451	0.207	16.628	0.000	0.11
	Temperature	-0.033	0.009	-3.574	0.000	
	Time	-0.008	0.004	-2.067	0.039	
Cyanobacteria	Intercept	2.367	0.402	5.880	0.000	0.08
	Temperature	-0.063	0.018	-3.465	0.001	
	Time	0.006	0.008	0.839	0.402	

563



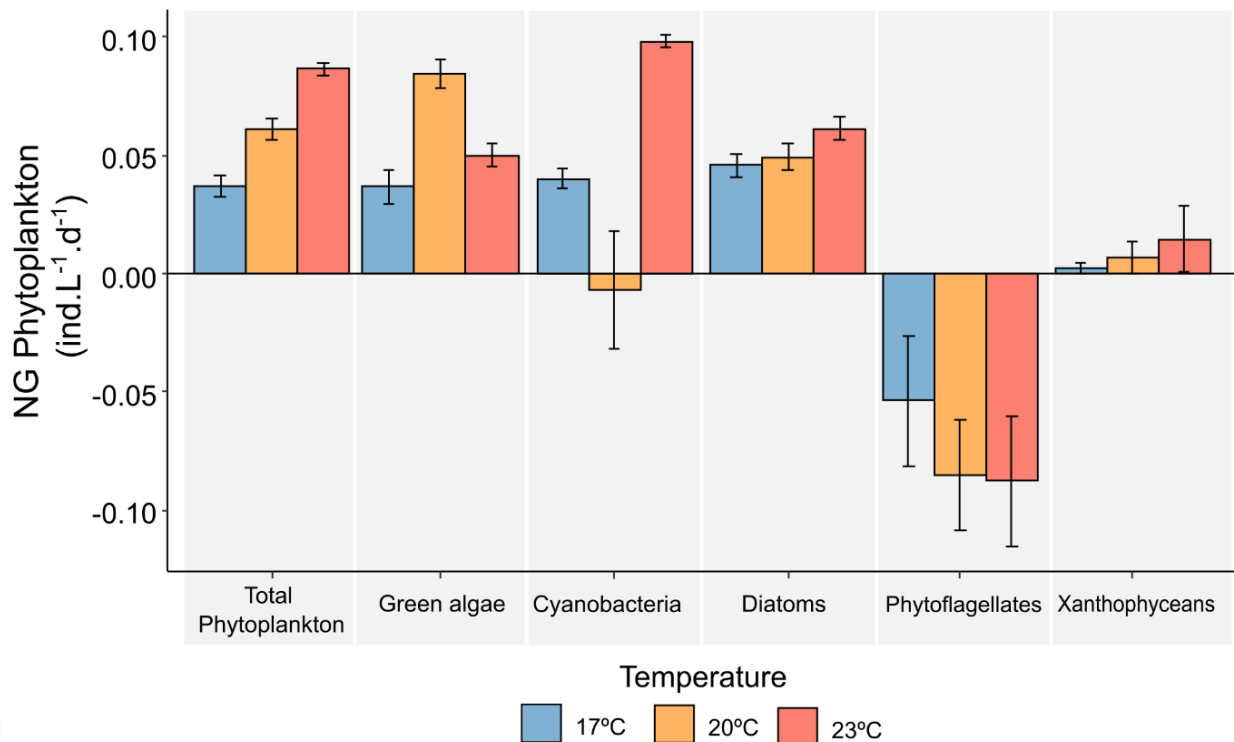
564

565 **Figure 4:** Variation of phytoplankton biomass (as biovolume) with temperature through
 566 the 32 days of the experiment: total (a), green algae (b), cyanobacteria (c). The dotted
 567 line indicates that on the 16th day of the experiment the final temperatures were reached.
 568 The central point denotes the mean value and, whiskers represent standard error for each
 569 temperature in each experiment day.

570

571 **Table 2.** Generalized linear models, indicating regression estimate, standard errors (SE),
 572 z-value, and P-values of models predicting, as from the 16th day, of the total
 573 phytoplankton biomass, green algae, and cyanobacteria. Significant results are shown in
 574 bold.

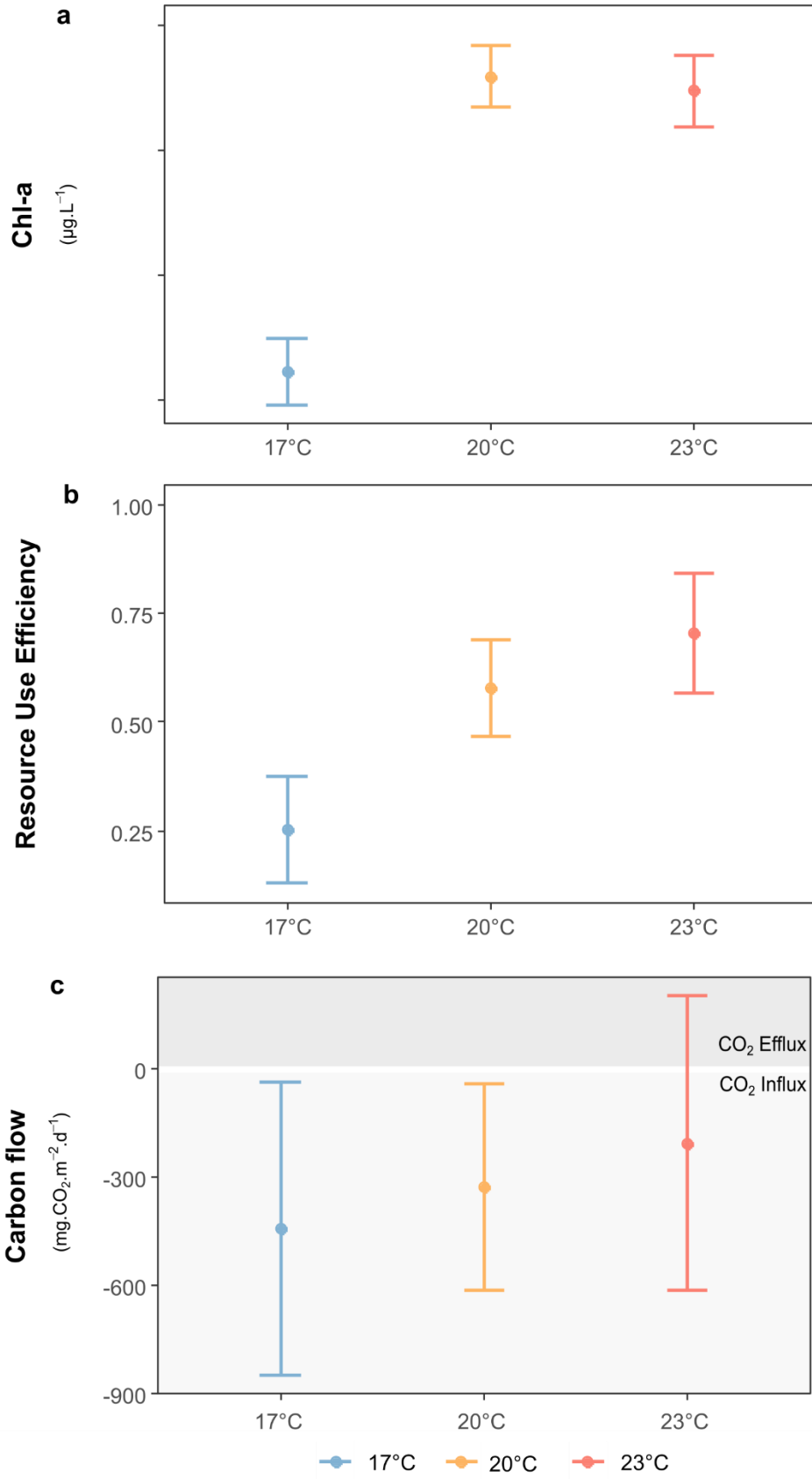
Response variable	Predictors variables	Estimate	SE	z-value	P	Pseudo R ²
Total Biomass	Intercept	-78.92	9.375	-8.417	0.000	0.50
	Temperature	3.562	0.416	8.559	0.000	0.77
	Time	1.489	0.177	8.385	0.000	0.78
Green algae	Intercept	-0.429	6.530	-0.066	0.948	0.03
	Temperature	0.197	0.290	0.679	0.499	0.579
	Time	0.278	0.124	2.248	0.026	0.580
Cyanobacteria	Intercept	-78.77	9.929	-7.934	0.000	0.41
	Temperature	3.235	0.441	7.338	0.000	0.581
	Time	1.240	0.188	6.592	0.000	0.582



583

584 **Figure 5.** Net growth rate (NG) of phytoplankton density of total and main
 585 phytoplankton groups under the three temperature treatments. The bars plot denotes
 586 mean and, whiskers represent standard error.

587



588

589 **Figure 6.** Different ecosystem responses to warming: variation of chlorophyll-a
590 concentration (Chl-a - log-transformed) (a), resource use efficiency (log-transformed)
591 (b), and carbon flow (c). Positive values of carbon flow indicate net CO₂ emissions
592 while negative values indicate net CO₂ sequestration by the phytoplankton communities
593 in three treatments. The central point denotes the mean value and, whiskers represent
594 standard error for each temperature in each experiment day.

595 2.3.2 CO₂ fluxes

596

597 Although we registered high CO₂ influx values in the experiment (i.e. CO₂
598 retention in the aquatic environment, due to net photosynthesis), the ecosystem
599 metabolic balance changed with temperature increase. The magnitude of influx
600 decreased with warming, almost transforming the microcosms in sources of CO₂ to the
601 atmosphere (i.e., CO₂ liberation to the atmosphere, net respiration). These results
602 support our hypothesis of positive feedback between blooms and warming, through an
603 increase in CO₂ efflux to the atmosphere.

604 The highest daily CO₂ efflux (580.24 CO₂.mg⁻². d⁻¹) occurred on the 10th day of
605 the experiment in the 23°C treatment, and the highest CO₂ influx (-1579.58 CO₂.mg⁻². d⁻¹)
606 occurred on the 25th day of the experiment in the 17 °C treatment (Fig. 8A). We
607 found significant differences in CO₂ concentration between the temperature treatments
608 (F-value (2, 80) = 4.54; P = 0.013), the differences being the largest between 17°C and
609 23°C (Tukey test p = 0.027). The highest CO₂ mean values (i.e., the lowest influx, -184
610 CO₂.mg⁻². d⁻¹) occurred in the 23°C treatment while the lowest (i.e., the highest influx, -
611 445 CO₂.mg⁻² d⁻¹) was registered in the 17°C treatment (Fig.6c).

612 2.4 Discussion

613

614 Climate warming and eutrophication are two major challenges for lacustrine
615 management worldwide (Carter & Schindler, 2012; Moe et al., 2016). Our results show

616 experimentally that, under future scenarios of climate warming, the phytoplankton
617 community composition can respond strongly, affecting ecosystem functions such as
618 biomass production, resource use efficiency, carbon flux balance.

619 We conducted our experiment at the community level, with a large initial species
620 pool that represented high genetic, taxonomic, and functional variability. Warming may
621 promote deep changes in the organization and diversity of the phytoplankton
622 community. We found a negative effect of temperature increase on phytoplankton
623 richness, with a functional change in the structure of the phytoplankton community and
624 dominance patterns through time, with the substitution of eukaryotic by cyanobacteria
625 and biotic homogenization, especially in the treatment with the highest temperature.

626 Besides composition changes, phytoplankton biomass increased with increasing
627 temperature, especially green algae and cyanobacteria. In line with Verbeek et al.
628 (2018) showing that climate change favors species with wider temperature ranges or
629 higher temperature optima, such as green algae and cyanobacteria species in an
630 experimental study with similar time duration to ours. Chlorophyta and Cyanobacteria
631 are both considered fast-growing r-strategists that can endure higher temperatures and
632 are favored in changing environments by their fast turnover, thus being able to develop
633 blooms in a very short period.

634 Some studies have shown that green algae also have their optimal growth at high
635 temperatures (De Senerpont Domis et al., 2007; Low-Décarie et al., 2011; Lürding et al.,
636 2013). The latter experimentally found that some chlorophyceans (some of them also
637 registered in this study, as *Desmodesmus* spp.) presented their optimal growth at high
638 temperatures (approximately 20 °C), with no significant difference in the optimal
639 growth temperature between cyanobacteria and green algae. Our results indicate that a
640 community approach is needed to detect changes in dominance patterns. However, due

641 to the competitive advantage of Cyanobacteria to rapidly sequester nutrients, they can
642 grow faster and outcompete other algae that are less efficient in nutrients uptake in the
643 warming environment (Rasconi et al., 2017).

644 *Raphidiopsis raciborskii* (CyanoHABs), of a tropical origin but invasive over a
645 wide distribution in temperate zones (Wiedner et al., 2007), was the main responsible
646 for the highest biomass of cyanobacteria in the highest temperature conditions. This
647 result corroborates other studies that showed that species of this group are favored with
648 warming (Kosten et al., 2011; Bonilla et al., 2016; Rasconi et al., 2017; Huisman et al.,
649 2018; Gray et al., 2019; Ho et al., 2019). This species may also use limiting resources
650 more efficiently than other cyanobacteria species due to its high affinity for and high
651 storage capacity of P (Isvánovics et al., 2000; Wu et al., 2012). To maintain several
652 ecosystem functions multiple species with different traits are necessary (Hector &
653 Bagchi, 2007; Mouillot et al., 2013a; Litchman et al., 2015). Thus, a decrease in
654 diversity (as a consequence of the dominance of traits) has been associated with a
655 general decrease in ecosystem functionality (Gamfeldt et al., 2008, 2013).

656 We found greater resource use efficiency with higher temperatures allow
657 phytoplankton to yield a higher carbon-based biomass per unit cellular P. This suggests
658 that RUE of phytoplankton increases with temperature, confirming earlier findings
659 showing that primary productivity increases with temperature (Kerkhoff et al., 2005;
660 Lovelock et al., 2007; De Senerpont Domis et al., 2014; Verbeek et al., 2018).

661 Ecologists have obtained inconsistent conclusions when analyzing the influence
662 of phytoplankton diversity on RUE of phytoplankton, as this relationship depends on
663 the composition of the community and species-specific RUE (Tian et al., 2017; Verbeek
664 et al., 2018). Previous studies supported the existence of a positive relationship between
665 species richness and high community RUE values (Ptacnik et al., 2008; Striebel et al.,

2009; Chai et al., 2020), as a larger number of coexisting species exploit a wider range of niches and can use limiting resources more efficiently (Tilman et al., 1997). However, our results showed an opposite relationship, with higher RUE in the communities dominated by few cyanobacteria genera which are more efficient in resource use and limiting the growth of other species (Roy & Chattopadhyay, 2007; O’Neil et al., 2012; Filstrup et al., 2014a; Sukenik et al., 2015). Lost diversity may be related to the fact that an isolated ecosystem has no chance of recovering species through regional diversity through immigration (Hillebrand et al., 2010).

Despite the higher maximum biomass of the phytoplankton community at higher temperatures, the low nutritional quality nutrient of these primary producers may negatively affect higher trophic levels (Soares et al., 2009). This may generate cascading effects on the superior trophic levels due to the reduced nutritional quality of cyanobacteria (Hassett et al., 1997). Phytoplankton is an important determinant of water quality and is key food production for heterotrophs, supporting the fish stocks in aquatic systems. Climatic effects on these autotrophs are thus of considerable interest. (Filstrup et al., 2014) for example, found lower zooplankton biomass in ecosystems dominated by a few phytoplankton genera. Specifically, the biomass stored in non-edible autotrophs (in this case, *Raphidiopsis raciborskii*, which has filaments that obstruct filtrating appendices of zooplankton species) alters the efficiency of resource transference to herbivores and the structure of trophic interactions.

Our results support previous investigations showing that warmer waters promote the dominance of cyanobacteria in phytoplankton communities in different approaches: experimental data with cultures (Staehr & Birkeland, 2006); experimental data with natural communities (De Senerpont Domis et al., 2014; Verbeek et al., 2018; Machado et al., 2019); long-term experimental data (Burgmer & Hillebrand, 2011; Yvon-

691 Durocher et al., 2011; Rasconi et al., 2017); observational studies (Paerl & Huisman,
692 2008; Kosten et al., 2012; Beaulieu et al., 2013); and paleolimnological studies (Pal et
693 al., 2015). On the other hand, studies show that the correlation of cyanobacteria biomass
694 and temperature in lakes is still unclear, depending on the lake trophic state or trophic
695 interactions (e.g. resistance to grazing) (Lürling et al., 2013; Kraemer et al., 2017;
696 Almanza et al., 2019; Gerhard et al., 2019). We designed our experiment to focus on
697 increasing mean temperature as a driver of changes in the phytoplankton community,
698 providing non-limiting nutrients and no predation conditions, to avoid indirect or
699 interacting effects of warming. Therefore, we can assume that warming directly drove
700 cyanobacteria proliferation.

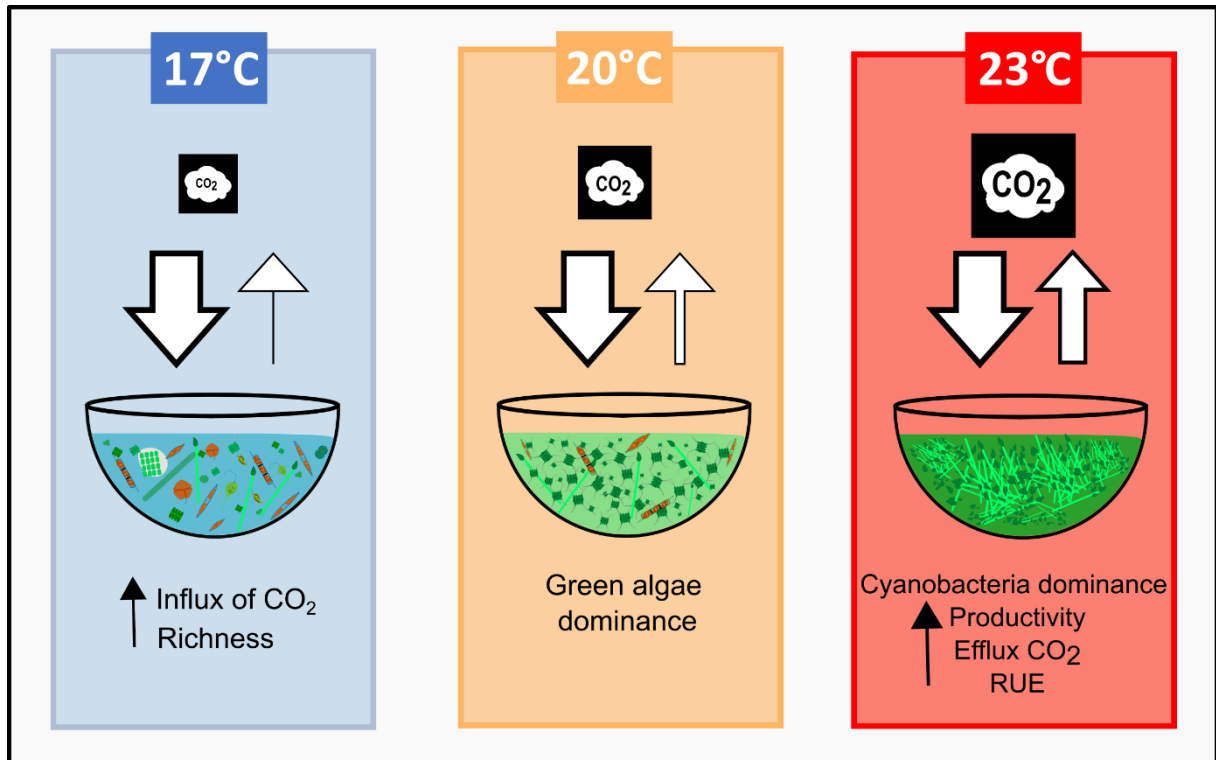
701 Supporting our hypothesis, we found evidence of positive feedback between
702 eutrophication and warming through changes in carbon emissions. Although all our
703 treatments can be considered as net carbon sinks, the CO₂ sequestration in the
704 communities with higher temperatures was severely diminished. Therefore, warming
705 may increase CO₂ emissions from eutrophic lakes to the atmosphere (further
706 aggravating the greenhouse effect) and reduce the role of these aquatic systems as
707 carbon sinks (Yvon-Durocher et al., 2010, 2017). Our results are following studies
708 realized in a variety of natural ecosystems that highlight the strong dependence between
709 carbon efflux and temperature (Whiting & Chanton, 1993; Christensen et al., 2003;
710 Gedney et al., 2004). A recent study suggests that there are feedbacks (“Vicious loop”)
711 among cyanobacteria blooms occurrence, lake eutrophication, and climate warming
712 (Yan et al., 2017).

713 Besides, after the collapse of cyanobacteria, phytoplankton loss processes, as
714 sedimentation and decomposition, are intensified. Many studies have suggested that
715 organic carbon in sediments is mineralized into gas emissions of greenhouse gas, such

716 as CO₂ (Gudasz et al., 2010; Bastviken et al., 2011; Marotta et al., 2014). Furthermore,
717 we found that the warmest treatment caused a decrease in dissolved oxygen
718 concentration. This may be related to lower oxygen solubility and/or to an increase in
719 the metabolic rates of the organisms at higher temperatures (Diaz & Breitburg, 2009). A
720 lower dissolved oxygen concentration may, in turn, promote a stronger release of
721 methane and nitrous oxide, further promoting warming. Besides, intensifies the
722 respiration of planktonic organisms (Yvon-Durocher et al., 2015) and decomposition
723 rates (Geraldés et al., 2012) both of which are processes that involve oxygen
724 consumption.

725 Experiments studies are a fundamental tool to predict how scenarios of climate
726 change will affect processes at the ecosystem levels and allow us to make predictions
727 about how organisms may respond to changing environments (Stewart et al., 2013),
728 despite their obvious limitations (Benton et al., 2007). In our case, the microcosms did
729 not mimic a catchment context and, consequently, received no terrestrial organic
730 carbon. Benthic processes and even pelagic trophic relations were purposely excluded.
731 We used a highly diverse initial community and isolated it from potentially confounding
732 factors. Thus, we were able to capture a pure pelagic response to warming, exclusive to
733 the phytoplankton community, and how such responses translated to ecosystem
734 functions (e.g., biomass production, energy transfer, and carbon cycle). Warming may
735 aggravate the negative effects of eutrophication through the enhancement of
736 cyanobacteria. Here, we also found experimental evidence of positive feedback between
737 eutrophication symptoms (blooms) and warming, via higher CO₂ emission rates in
738 cyanobacteria dominated warmer systems (Fig.7).

739



740

741 **Fig.7.** Under climate warming scenarios, ecosystem functions are affected. For example, with
 742 warming, environments become more productive, due to eutrophication with the dominance
 743 of cyanobacteria (less diversity), and greater resource use efficiency (RUE). This scenario
 744 gradually changes the ecosystem's metabolic balance changed, taking the lakes in the
 745 direction of being CO₂ sources for the atmosphere. We evidenced positive feedback
 746 mechanisms relationship among climate warming, lake eutrophication, and cyanobacteria
 747 blooms (vicious loop).

748

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- 1066

1067 **3 WARMING PROMOTES RESILIENCE OF CYANOBACTERIAL BLOOMS TO**
1068 **SIMULATED RAINFALL EXTREME EVENTS**

1069

1070

1071 **ABSTRACT**

1072

1073 Cumulative stressors including climate change (warming, extremes rainfall events)
1074 and eutrophication have increased the frequency and severity of cyanobacteria blooms. In this
1075 study, we conducted an indoor short-term experiment to test how natural phytoplankton
1076 communities reacted to an extreme precipitation event under different temperature scenarios.
1077 Our main hypothesis was that communities less stressed by warming would be more resistant
1078 and would more rapidly recover their prior biomass and resource use efficient (RUE), than
1079 communities under higher temperature. Contrary to our hypothesis, in the communities
1080 exposed to lower temperature (supposedly less stressed), the effects of the disturbance were
1081 more evident. Warming promoted cyanobacteria-dominated communities showed better
1082 recovery of biomass and a thus higher ability to withstand the changes caused by the extreme
1083 rainfall event. The blooms, dominated by *Raphidiopsis raciborskii*, returned in 10 days to
1084 similar values of RUE and chlorophyll-a compared to those not subjected to the simulated
1085 rainfall event. Phytoplankton communities developed under lower temperatures did not show
1086 such resilience. Our results highlight the increasing vulnerability of freshwater ecosystems to
1087 warming as well as to rainfall extreme events. Climate change clearly aggravates the negative
1088 effects of eutrophication through the enhancement of cyanobacteria, which, as demonstrated
1089 here, has great stability and recoverability after disturbances.

1090

1091 **Keywords:** Climate change, multiple stressors, eutrophication, resilience, resistance, rainfall.

1092 3.1 Introduction

1093 Climate change is often studied as a single stressor (most typically the increase in
1094 temperature) impacting the natural environment, although climate change also increases the
1095 frequency and magnitude of extreme events as well as overall patterns of precipitation (Field
1096 *et al.* 2014, Lehmann *et al.* 2015). In lakes, one of the most clear and expected outcomes of
1097 climate warming is the positive interaction with eutrophication process and symptoms,
1098 leading to an increase in the distribution and intensity of cyanobacterial blooms across the
1099 globe (Garcia-Pichel *et al.*, 2003; Paerl & Otten, 2013; Harke *et al.*, 2016; Paerl *et al.*, 2020;
1100 Weber *et al.*, 2020), and the consequent major threat to freshwater quality and global water
1101 security (Codd *et al.*, 2005; Michalak, 2016; Richardson *et al.*, 2018; Jeppesen *et al.*, 2021).
1102 Although nutrients seem to be the more important predictor of cyanobacterial biovolume, as
1103 lakes become more eutrophic cyanobacteria become more sensitive to the interaction of
1104 nutrients and temperature (Rigosi *et al.*, 2014).

1105 Associated with warming, extreme rainfall events can have multiple effects and the
1106 mechanism behind alterations in precipitation and its impacts on blooms is not well
1107 understood. The responses of phytoplankton communities to disturbances are influenced by
1108 many factors, including the type of aquatic environment (reservoirs, shallow lakes, and deep
1109 lakes) (Doubek *et al.*, 2017; Hayes *et al.*, 2017; Richardson *et al.*, 2018), abiotic and biotic
1110 conditions, and extant phytoplankton community composition (Stockwell *et al.*, 2020). Thus,
1111 changes in precipitation patterns can have multiple effects at different scales and levels, with
1112 different direct and indirect impacts on phytoplankton and particularly on cyanobacteria.

1113 For instance, future increases in mean precipitation may promote the occurrence of
1114 cyanobacteria, due to higher nutrient input by increased runoff from the catchments
1115 (Ockenden *et al.*, 2017), thus increasing blooms (Shaw *et al.*, 2001; Jeppesen *et al.*, 2011). A
1116 long-term reduction in mean precipitation can also favor cyanobacteria due to a higher

1117 stratification of the water column and higher concentration of nutrients (Brasil et al., 2016).
1118 This may occur when extreme rainfall is followed by periods of droughts (Paerl & Huisman,
1119 2009; Havens & Ji, 2018). Short-term changes, such as intense rainfall events can generate
1120 unfavorable conditions for cyanobacteria due to dilution and flushing; causing either a
1121 decrease in biomass or a complete collapse of the bloom due to destratification of the water
1122 column (Reynolds et al., 2012; Sadro & Melack, 2012). Indirectly, increased flow may also
1123 result in environmental changes and consequently affect biological responses, such as changes
1124 in selection pressures that affect community composition and diversity (James et al., 2008;
1125 Reichwaldt & Ghadouani, 2012). The mechanisms determining ecosystem response and
1126 recovery to climate extremes remain unclear, making vulnerability assessments uncertain
1127 (Kayler et al., 2015; De Boeck et al., 2018).

1128 Cumulative stressors linked to climate change, as warming, anthropogenic
1129 eutrophication, extreme weather, all factors associated with promoting cyanobacterial blooms
1130 put at risk the functioning in freshwater ecosystems (e.g. resource use efficiency – RUE)
1131 (Wagner & Adrian, 2009; Kosten et al., 2012; Filstrup et al., 2014). In addition to the
1132 intensification of multiple environmental stressors, many aspects of global change are
1133 expected to alter the frequency, variance and timing of disturbances (De Laender et al., 2016;
1134 Donohue et al., 2016; Radchuk et al., 2019). After a major disturbance, the re-establishment
1135 of species is highly variable and depends on the extent of physical alteration of the
1136 environment, species growth rates, competition, predation and other factors (Ji et al., 2018).

1137 Ecological ‘stability’ is the core concept describing potential responses to such
1138 changes, a concept of central importance for understanding present-day and predicting future
1139 ecosystem dynamics. The multifaceted concept of ecosystem stability includes: resistance, as
1140 the ability to withstand the perturbation (Pimm, 1984; Donohue et al., 2016), recovery, the

1141 ability to return to their pre-disturbance state, and resilience, the time needed to return to their
1142 pre-disturbance state (Holling, 1973; Orians, 1975).

1143 Understanding the ecosystem responses and stability against long-term and short-term
1144 stressors is crucial to be able to suggest better management and restoration strategies
1145 (Scheffer et al., 2001; Pecl et al., 2017; De Boeck et al., 2018). In this sense more diverse
1146 systems are expected to have greater ecosystem stability and greater resistance to a
1147 disturbance, reflecting interspecific complementarity, higher resource use efficiency (Tilman
1148 et al., 2014). Still, it is expected that eutrophic systems, due to the dominance of a single
1149 species and less diversity, may have a lower level of ecosystem stability in response to
1150 climate change (Filiz et al., 2020).

1151 Here, we conducted an indoor short-term experiment to test how distinct natural
1152 phytoplankton communities respond to climate change (a disturbance, i.e., a simulated rainfall
1153 extreme event and a stress, i.e., different warming levels). We expected that communities less
1154 stressed by warming would be more resistant against a disturbance and would have the ability
1155 to recover their pre-disturbance biomass and functioning faster than the other phytoplankton
1156 communities.

1157 **3.2 Methods**

1158 3.2.1 Experimental Design

1159 We conducted a short-term indoor controlled microcosm experiment in a laboratory
1160 located at the *campus* of the Universidad de la Republica (UdelaR), Maldonado, Uruguay (34°
1161 54'53'S e 54 ° 56'31 " W '), between July 23 and August 2 of 2019 (southern hemisphere
1162 winter). We compared the response to a disturbance by three distinct phytoplankton
1163 assemblages, which resulted after four weeks of the same original community being subject to
1164 three temperature treatments. The control treatment, (i) 17 °C, corresponded to the annual
1165 mean temperature of the aquatic ecosystems in the same region where the experiment was

1166 conducted (Pacheco et al., 2010). The second treatment, (ii) 20 °C, represented a scenario
1167 with an increase of 3 °C in relation to the annual mean temperature; and the third treatment,
1168 (iii) 23 °C, represented an increase of 6 °C in relation to the mean temperature (Fig.1). Room
1169 temperature was manipulated in all treatments using electronic temperature controllers
1170 (Eliwell ID Plus 961), including the control treatment.

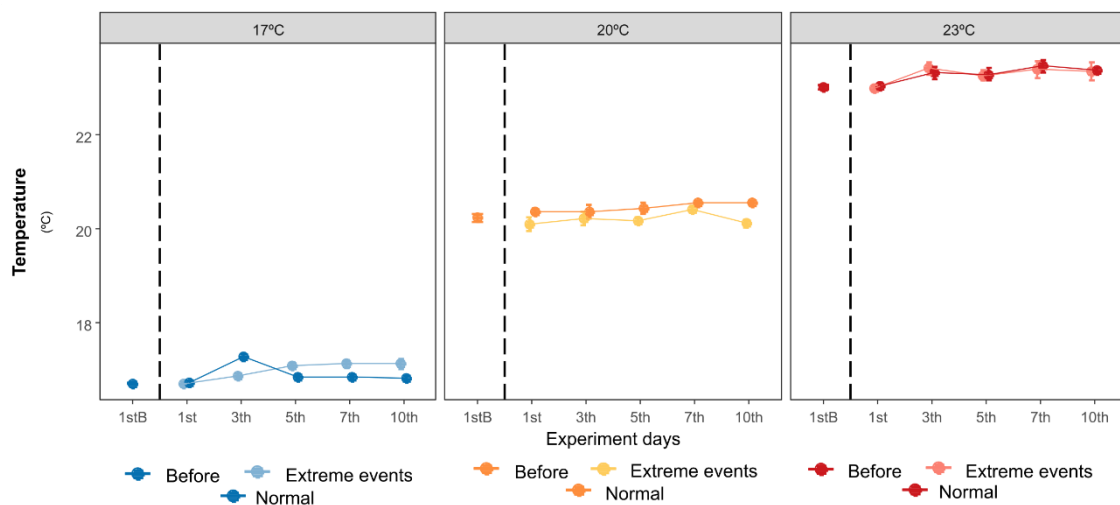
1171 The microcosm was cylindrical polyethylene 5-L aquarium (23.5 cm of diameter and
1172 10 cm of height). We installed water circulation pumps (model Resun AC-9903) at the bottom
1173 of each aquarium to avoid phytoplankton sedimentation (Flury et al., 2010; Zhang et al.,
1174 2015). Besides, we moved the water manually and delicately twice a day with a stick. Light
1175 regime consisted of circa 80 lum/ft² of 12/12 hs ratio of light/darkness generated by
1176 fluorescent lamps with a light spectrum similar to that of the sun. All experimental units were
1177 placed at the same height to homogenize temperature and incident light. We randomly
1178 distributed twenty-four microcosms between the three temperature treatments (i.e., n=8
1179 replicates each).

1180 To prepare the natural phytoplankton original inoculum for the experiment, we
1181 obtained water samples at the subsurface of the limnetic region in a series of subtropical
1182 shallow lakes situated at the Uruguayan coastline (34° S 56° W). These lakes comprised a
1183 trophic gradient from mesotrophy to hypertrophy (Kruk et al., 2009). We collected
1184 phytoplankton using a plankton net with a mesh size of 20 µm to remove most zooplankton
1185 and added some non-filtered water to include smaller species (< 20 µm). We used this
1186 strategy to maximize the sampling of less abundant species and guarantee that most
1187 taxonomic groups were present in the experimental units. We distributed 0.5 L of
1188 concentrated inoculum for each of the 24 experimental units and added 2.5 L of dechlorinated
1189 water enriched with a nutrient solution to reach a final volume of 3 L.

1190 After 4 weeks, three distinct communities emerged (summarizing 72 taxa,
 1191 Supplementary Figure 1). The lowest mean values of phytoplankton biomass were recorded in
 1192 the 17°C treatment ($13 \text{ mm}^3.\text{L}^{-1} \pm 0.66 \text{ mm}^3.\text{L}^{-1}$) and the highest mean biomass values for
 1193 20°C treatment ($23.96 \text{ mm}^3.\text{L}^{-1} \pm 1.69 \text{ mm}^3.\text{L}^{-1}$) and 23°C treatment ($24.33 \text{ mm}^3.\text{L}^{-1} \pm 2.27$
 1194 $\text{mm}^3.\text{L}^{-1}$). Cyanobacteria (as *Raphidiopsis raciborskii* (Wołoszyńska) Aguilera, Berrendero
 1195 Gómez, Kastovsky, Echenique & Salerno, before *Cylindrospermopsis*) was the main group
 1196 for the treatments of 17°C and 23°C with mean values of $7.35 \text{ mm}^3.\text{L}^{-1}$ and $18 \text{ mm}^3.\text{L}^{-1}$,
 1197 respectively whereas green algae (as *Desmodesmus magnus* (Meyen) Tsarenko and
 1198 *Staurastrum* sp.) was the largest contributor under 20°C ($12 \text{ mm}^3.\text{L}^{-1}$). For more details on the
 1199 methodology see Moresco et al. (*in press*)

1200 To simulate extreme rainfall events four experimental units from each temperature
 1201 were randomly selected, and, after intense homogenization, 1 L of water was removed and 1
 1202 L dechlorinated water enriched with a nutrient solution was added. The experiment lasted for
 1203 10 days, as the short life cycle of phytoplankton allows the reproduction of several
 1204 generations in a few days (Reynolds, 2006).

1205



1206

1207 **Figure 1.** Mean daily temperature values in each temperature treatment. Light colors
 1208 correspond to the extreme rainfall treatment and dark colors non disturbed treatments. The

1209 dotted line indicates the occurrence of the extreme rainfall event (1stB). The central point
1210 denotes mean value and whiskers represent standard error.

1211 3.2.2 Monitoring of limnological variables

1212

1213 In all microcosms, every two days we measured water temperature ($^{\circ}\text{C}$), pH, and
1214 electric conductivity using a HANNA multiparametric probe, and dissolved oxygen (mg L^{-1})
1215 and oxygen saturation (%) using an Oxyguard Handy Polaris. We sampled water on the first
1216 and the last day (10th) to determine total phosphorous (TP; $\mu\text{g L}^{-1}$), total nitrogen (TN; $\mu\text{g L}^{-1}$)
1217 ¹), reactive soluble phosphorous (SRP; $\mu\text{g L}^{-1}$), nitrate (N-NO₃; $\mu\text{g L}^{-1}$), and ammonia (N-
1218 NH₄; $\mu\text{g L}^{-1}$), as well as chlorophyll-a, according to APHA (2005). Phytoplankton was
1219 maintained in a medium following the Redfield ratio. Weekly, a nutrient solution was added
1220 to each microcosm to compensate for the losses due to evaporation as to reach the initial
1221 volume of 3 L (McKee et al., 2000; Ekvall & Hansson, 2012). The mean and standard error of
1222 nutrients in microcosms were $2431.9 \pm 176.2 \mu\text{g L}^{-1}$ of TN and $81.2 \pm 2.3 \mu\text{g L}^{-1}$ of TP.

1223

1224 3.2.3 Phytoplankton

1225 We sampled phytoplankton every five days (day 1 before, and after the extreme
1226 rainfall event, day 5 and day 10). Before taking each sample, we homogenized the water
1227 manually to avoid missing any species due to sedimentation. We sampled phytoplankton
1228 directly with flasks and fixed them immediately with acetic Lugol solution. Counting of
1229 individuals (cells, colonies, and filaments) of phytoplankton followed the Utermöhl method
1230 (Utermöhl, 1958) and Lund *et al.* (1958). The biomass ($\text{mm}^3 \cdot \text{L}^{-1}$) was considered as
1231 biovolume, which was estimated by multiplying the density of each taxon by its volume. We
1232 estimated the cell volume by calculating the volume of the geometric shape that was the most
1233 similar to each cell form (Sun & Liu, 2003). We also estimated the community resource use
1234 efficiency (thereafter RUE), defined as the ratio between the phytoplankton biomass

1235 production and TP, as a proxy for ecosystem productivity (Ptacnik et al., 2008; Olli et al.,
1236 2015; Verbeek et al., 2018).

1237

1238 3.2.4 Data analyses

1239

1240 To verify the existence of significant differences between the occurrence or absence of
1241 the extreme rainfall events along time we performed two permutational multivariate analyses
1242 of variance (PERMANOVA) (Anderson, 2001b), using as response variables the biomass of
1243 different phytoplankton species. The variation in the biomass trajectory of the phytoplankton
1244 community and its capacity to return in the initial stage (i.e., resilience) was visualized using a
1245 nonmetric multidimensional scaling (NMDS). In NMDS, distances were calculated using the
1246 Bray–Curtis similarity index with the resolution distortion in two dimensions expressed by the
1247 value S (stress) (Clarke, 1993). Resistance and recovery were calculated for phytoplankton
1248 total biomass following (Hillebrand et al., 2018). Resistance and recovery were calculated
1249 with phytoplankton total biomass sampling following this ratio: $\ln(\text{disturbed treatment/un}$
1250 $\text{disturbed treatment})$. We use initial biomass (after disturbance) for the calculation of the
1251 resistance and the final biomass for the recovery. We ran all analyses in software R version
1252 3.3.2 (R Development Core Team, 2021) , using packages “vegan” (Oksanen et al., 2018).

1253

1254 **3.3 Results**

1255

1256 The rainfall disturbance led to different responses by the three distinctive
1257 phytoplankton communities developed under the three temperature treatments. The
1258 disturbance affected abiotic variables analyzed (Supplementary Figure 2) and had significant
1259 effects in biomass at 17°C ($P = 0.008$) and 20°C ($P = 0.01$) and not at 23°C ($P = 0.607$),
1260 according to the PERMANOVA. The non-significant effect of disturbance highlights the

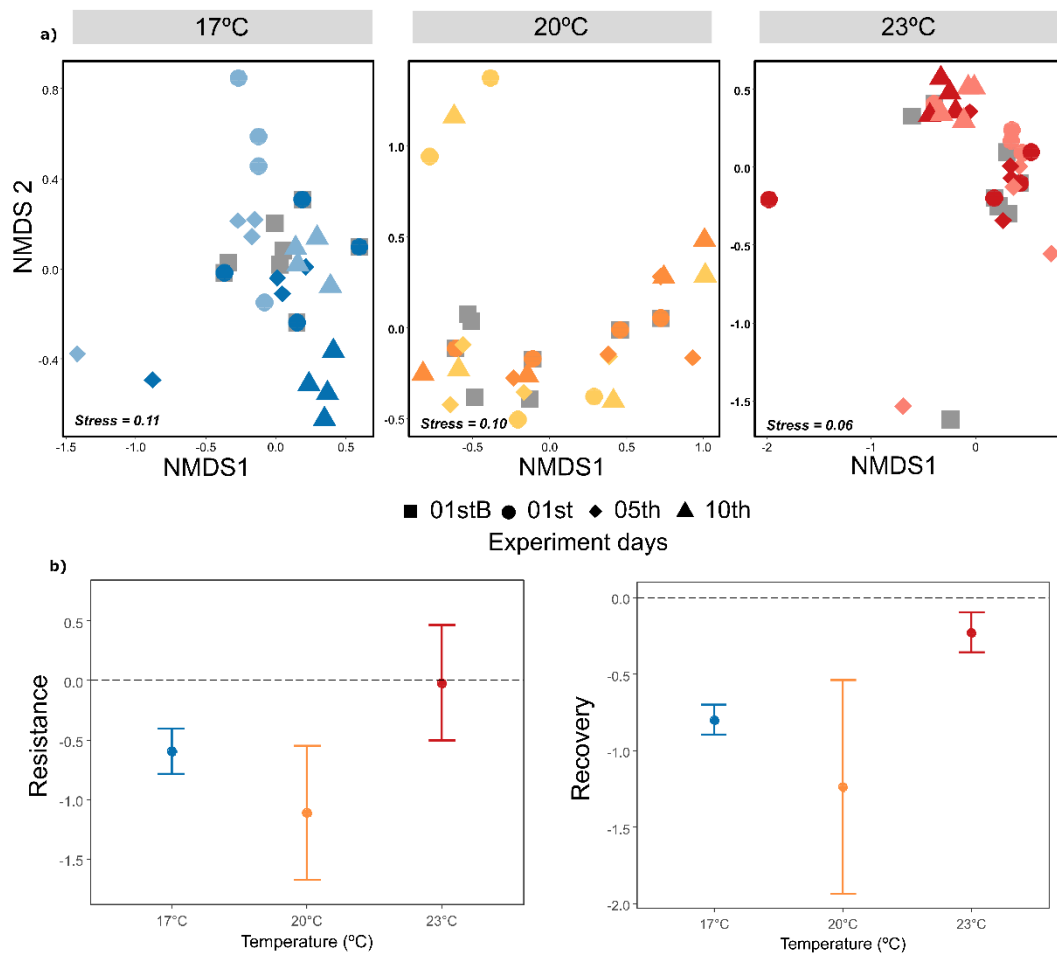
1261 resilience of the phytoplankton community developed under the warmest conditions (Table
 1262 1). In the phytoplankton communities developed under the lower temperatures (i.e.,
 1263 supposedly less stressed), the effects of the disturbance were more evident. Also, time had
 1264 significant effects on biomass ($P = 0.001$), except for 20°C (Table 1).

1265 **Table 1:** Effects of disturbance and time according on phytoplankton biomass under the three
 1266 temperature treatments, according to PERMANOVA test. Significant P values ($P < 0.05$) are
 1267 highlighted in bold.

Temperature	Df	17°C			20°C			23°C		
		R ²	F-value	P	R ²	F-value	P	R ²	F-value	P
Time	3	0.24	2.919	0.001	0.09	0.958	0.472	0.27	3.528	0.001
Disturbance	2	0.14	2.357	0.008	0.16	2.763	0.017	0.04	0.757	0.607

1268

1269 Changes in the trajectory of phytoplankton biomass over time and in relationship with
 1270 the disturbance indicate that the community in the warmest temperature has a greater capacity
 1271 for resilience, since on the last day of the experiment, even the environments that suffered
 1272 disturbance are close to those that are undisturbed. In contrast, less stressed environments had
 1273 a high dispersion of points, indicating less resilience (as shown in the NMDS) (Figure 2a).



1274

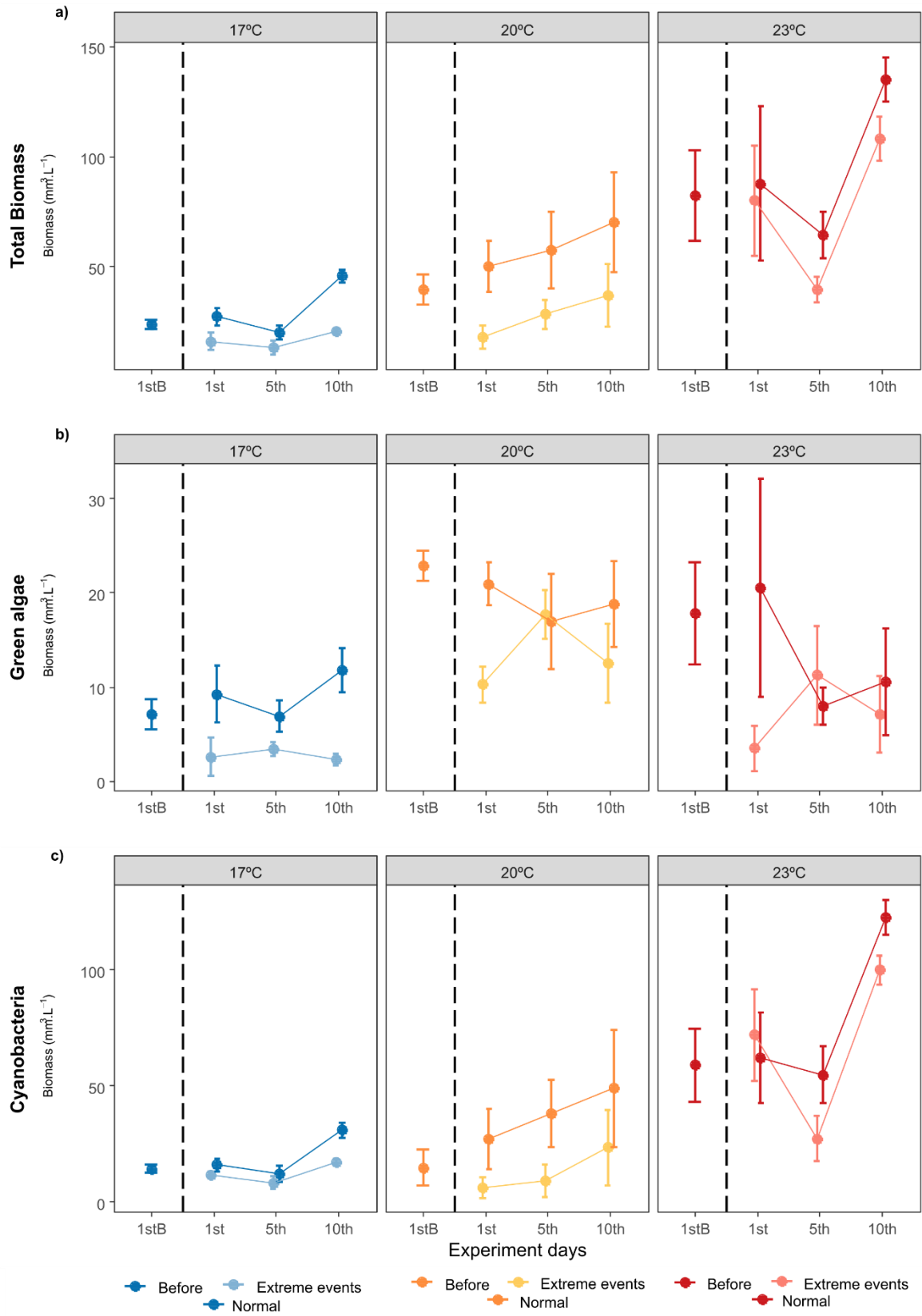
1275

1276 **Figure 2:** Trajectories (a), resistance (b), and recovery (c) of phytoplankton biomass. The
 1277 gray dots in the non-metric multidimensional scaling (NMDS) (a) indicate the period before
 1278 the extreme rainfall event (1stB), light colors indicate disturbed communities and dark colors
 1279 the undisturbed. Dashed lines in (b) and (c) are the benchmarks. The central point denotes the
 1280 mean value and whiskers represent standard error.

1281

In contrast to our hypothesis, phytoplankton communities developed under warming
 1282 conditions showed a better recovery of biomass and thus a higher ability to withstand the
 1283 changes caused by the extreme rainfall event simulated in our experiment. The less stress
 1284 communities (developed at low and intermediate temperatures) stayed distant from their
 1285 benchmark during the entire disturbance, indicating poor resistance and poor recovery. The
 1286 best recovery and higher resistance in the community developed under the warmest treatment

1287 was led by cyanobacteria, which on the last day of the experiment had a mean biomass of
1288 $122.09 \text{ mm}^3 \cdot \text{L}^{-1} \pm 7.32 \text{ mm}^3 \cdot \text{L}^{-1}$ in the control treatment and $99.02 \text{ mm}^3 \cdot \text{L}^{-1} \pm 6.15 \text{ mm}^3 \cdot \text{L}^{-1}$ in
1289 the disturbed treatment (Fig. 3, Supplementary Fig. 3), even surpassing the biomass achieved
1290 in the period before the disturbance ($63.86 \text{ mm}^3 \cdot \text{L}^{-1}$ Fig. 2a, Fig. 2b). Green algae did not
1291 recover their biomass at the levels before the disturbance at any temperature evaluated. The
1292 best performance for the group occurred at 20°C , even so with a difference of approximately
1293 $10 \text{ mm}^3 \cdot \text{L}^{-1}$ between the level before the disturbance and that of the extreme rainfall event
1294 (Fig. 3)

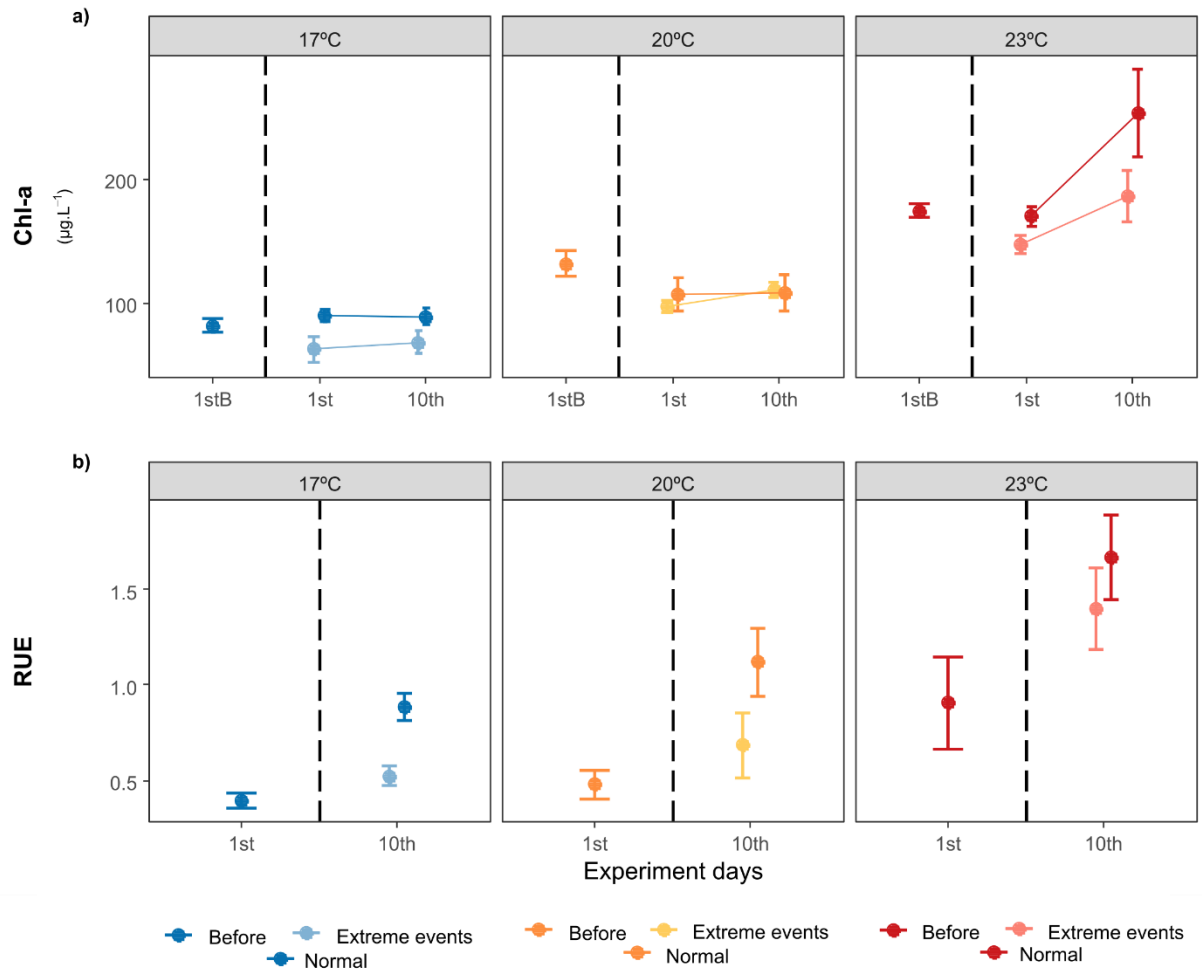


1296 **Figure 3:** Variation of phytoplankton biomass (biovolume) in the disturbance and
1297 temperature treatments: total (a), green algae (b), cyanobacteria (c). The dotted line indicates
1298 the occurrence of the extreme rainfall event (1stB). Light colors indicate communities subject
1299 to extreme rainfall and dark colors the undisturbed communities. The central point denotes the
1300 mean value and, whiskers represent standard error.

1301

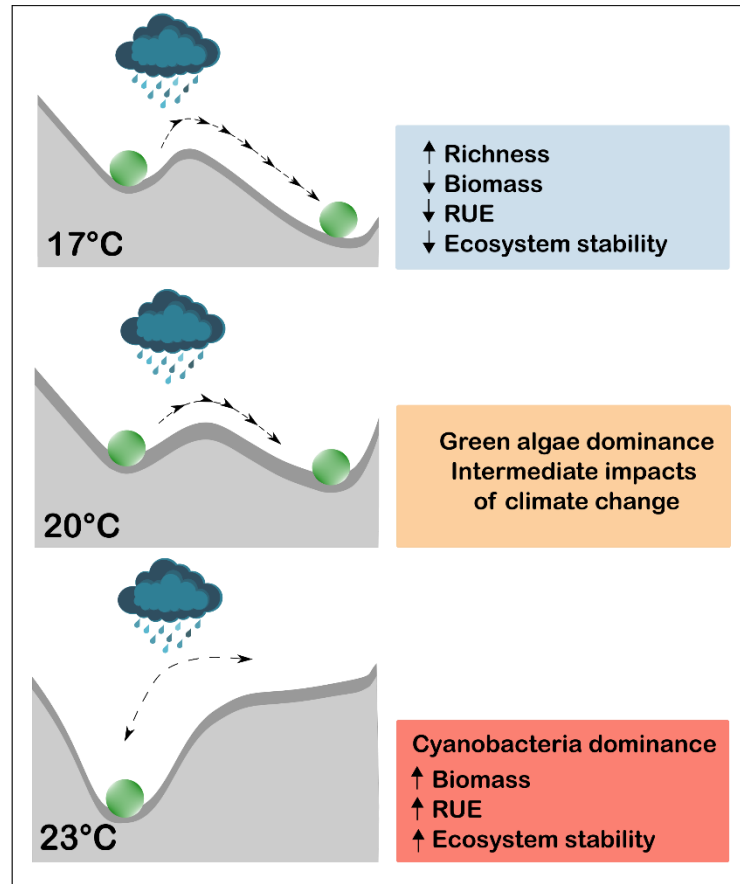
1302 The RUE and final chlorophyll-a concentrations were affected by the extreme rainfall
1303 event at the three temperatures evaluated. However, the disturbed communities under the
1304 warmest treatment (with higher biomass dominated by blooms of *Raphidiopsis raciborskii*)
1305 showed values of RUE and chlorophyll-a like the undisturbed communities. In the other
1306 temperature treatments, and particularly under the lowest temperature, the disturbed
1307 communities did not achieve values similar to the disturbed ones, except for chlorophyll-a in
1308 the 20°C (Fig. 4a, Fig. 4b). The conceptual diagram with the main results of our study is
1309 shown in Fig.5.

1310



1311

1312 **Figure 4.** Different ecosystem responses to warming: variation of chlorophyll-a concentration
 1313 (Chl-a) (a), resource use efficiency (RUE) (b) by the phytoplankton communities in three
 1314 treatments. Light colors indicate environments with the effect of extreme rainfall events and
 1315 dark colors without the effect. The central point denotes the mean value and, whiskers
 1316 represent standard error.



1317

1318 **Figure 5.** Under different experimental climate change scenarios, community structure and
 1319 ecosystem functions were affected. Different warming scenarios lead to different
 1320 phytoplankton communities (represent by ball), which had different resistance and resilience
 1321 against a short-term disturbance, as indicated by the biomass and resource use efficiency
 1322 (RUE) recovery.

1323

1324 3.4 Discussion

1325 We evaluated the effects of strong disturbance simulating an extreme climatic event,
 1326 on three phytoplankton communities pre-adapted to different levels of warming and under
 1327 non-limiting nutrient conditions. The study revealed that ecosystem stability and resilience
 1328 (recoverability of biomass and ecosystem function) depended on the dominant group.
 1329 Cyanobacteria-dominated communities, developed under the warmest treatment, were the
 1330 most resilient. Our results indicate that climate change, both via long-term stressors such as an

1331 increase in mean temperature and short-term perturbations such as intense rainfall events, may
1332 aggravate the negative effects of eutrophication through the enhancement of cyanobacteria.

1333 Freshwater ecosystems are becoming more vulnerable to multiple disturbances due to
1334 the combined and often synergistic effects of climate change and several human activities
1335 (Evtimova & Donohue, 2014; Hunter-Cevera et al., 2016; Paerl et al., 2016; Yang et al.,
1336 2017). Eutrophic environments seem particularly vulnerable to climate warming (Rigosi et
1337 al., 2014). In this sense, we found that phytoplankton communities developed under different
1338 temperatures respond differently to disturbances. Cyanobacteria blooms benefited and proved
1339 to be more resilient to a pulse disturbance than communities expectedly less stressed by
1340 warming, which in contrast, showed a slower response to the simulated extreme rainfall event.

1341 This indicates that the synergistic effect of multiple stressors is important in shaping
1342 the response of cyanobacteria and that a multiple factors approach could help better predict
1343 community responses to future environmental change. Several studies have shown that higher
1344 water temperature and rainfall regimes may play a key role in the proliferation of
1345 cyanobacteria, especially under eutrophic conditions (Bormans et al., 2005; Paerl & Huisman,
1346 2009; O'Neil et al., 2012; Wood et al., 2017).

1347 Our earlier results (Moresco et al., *in press*) had confirmed previous empirical findings
1348 that higher water temperatures promote cyanobacterial blooms (Paerl & Huisman, 2008;
1349 Kosten et al., 2012; Richardson et al., 2019), in our case through direct effects on
1350 phytoplankton community structure and performance. Many species of cyanobacteria may
1351 benefit in warmer conditions due to several traits (Carey et al., 2012; Mantzouki et al., 2016).

1352 Although the phytoplankton community is sensitive to losses due to extreme rainfall
1353 event (Reynolds et al., 2002; Carvalho et al., 2011; Stockwell et al., 2020), cyanobacteria can
1354 grow more efficiently at higher temperatures and persist longer under extreme wet/dry cycles
1355 (Paerl et al., 2016). Besides changes in the environmental conditions in aquatic systems

1356 caused by extreme rainfall events depends of others characteristics (hydrology of the
1357 catchment, waterbody type), eutrophic systems lead to higher biomass production
1358 (Reichwaldt & Ghadouani, 2012). This eutrophication process generally favors cyanobacterial
1359 blooms due to physiological characteristics that include the presence of aerotopes that allow
1360 them to move through the water column when there is thermal stratification, nitrogen-fixing
1361 capability, high affinity and phosphorus storage abilities, and the production of cysts
1362 (akinetes) (Weyhenmeyer et al., 2007; Salmaso et al., 2015). In addition, this process can be
1363 greatly intensified if after a large pulse of nutrients (via extreme rainfall event) followed by a
1364 dry period could benefit cyanobacteria ('perfect storm', Paerl *et al.* 2016). Thus, the survival
1365 chance of sensitive groups is much lower than that of adaptive groups (Filiz et al., 2020;
1366 Stockwell et al., 2020).

1367 Empirical and theoretical evidence suggests that more diverse assemblages show more
1368 stable productivity over time (Tilman, 1996; Isbell et al., 2015)). Although the environments
1369 with lower temperatures present greater species richness, they were not very resistant and had
1370 a low capacity to recover after the extreme rainfall event. A number of studies have shown a
1371 positive relationship between phytoplankton species richness and RUE which has been
1372 attributed to a more efficient use of resources in more species rich communities (Ptacnik et
1373 al., 2008; Striebel et al., 2009; Chai et al., 2020). However, we found that the RUE of
1374 phytoplankton increased with temperature, confirming earlier findings showing that primary
1375 productivity increases with temperature (Kerkhoff et al., 2005; Lovelock et al., 2007; De
1376 Senerpont Domis et al., 2014; Verbeek et al., 2018). Besides, after the dilution effect, the
1377 productivity of the ecosystem was only recovered in the cyanobacteria-dominated community
1378 developed under the warmest treatment. This result demonstrates that the most diverse
1379 communities do not necessarily better buffer the impacts of climate extremes on ecosystem
1380 functioning than less diverse communities. A reason for the faster recovery in the warmest

1381 conditions might be that the dominant *R. raciborskii* is able to use phosphorous more
1382 efficiently than other phytoplankton taxa, which concurs with studies finding strong effects of
1383 cyanobacteria on RUE in eutrophic lakes (Roy & Chattopadhyay, 2007; O’Neil et al., 2012;
1384 Filstrup et al., 2014; Sukenik et al., 2015). The temperatures used in our study were, however,
1385 in the range where no major differences in cyanobacterial or eukaryote algal growth rates
1386 were expected (Lürling et al., 2013).

1387 Experimental studies are a fundamental tool for understanding the complexity of how
1388 global climate change may impact freshwater ecosystems (Stewart et al., 2013), despite their
1389 obvious limitations (Benton et al., 2007). They allow developing a clearer mechanistic
1390 understanding of the interactions between multiple stressors, allowing quantification and
1391 comparison of individual stressor effects and their interactions (Crain et al., 2008; Piggott et
1392 al., 2015). In relation to limitations, in our case, the microcosms are isolated and, changes in
1393 the environmental conditions in aquatic systems caused by rainfall events will mainly depend
1394 on the hydrology characteristics of the catchment and waterbody, and the land use in area.
1395 Because benthic processes and even pelagic trophic interactions were purposely excluded, we
1396 were able to capture a pure pelagic response to the combined effects of climate change,
1397 exclusive to the phytoplankton community, and how such responses translated to ecosystem
1398 functions (e.g., biomass production, and ecosystem stability).

1399 We examined the response of phytoplankton communities subjected to stressors of
1400 climate changes, such as eutrophication, warming, and extreme rainfall events. The study
1401 revealed the increasing vulnerability of freshwater ecosystems phytoplankton community to
1402 warming as well as rainfall extreme events. Ecosystem stability in the warming microcosms
1403 was affected differently by the rainfall disturbance. Climate change stressors may aggravate
1404 the negative effects of eutrophication through the enhancement of cyanobacteria and, with
1405 increased frequency of extreme events, freshwater ecosystems may not have sufficient time to

1406 recover, and this could possibly lead to regime shifts. Besides, eutrophic and warm
1407 environments, dominated by cyanobacteria, showed greater stability and recoverability.

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1633 **4 FINAL CONSIDERATIONS**
1634

1635 In this thesis, we evaluated the effects of multiple factors related to climate change
1636 effects on natural freshwater phytoplankton community. Together with land use change,
1637 climate change is expectedly one of the major drivers threatening both aquatic biodiversity
1638 and several ecosystem services for the human population, such as water supply, recreation,
1639 culture, and irrigation. Thus, in the first paper, we tested how increasing temperatures
1640 influence natural phytoplankton diversity and CO₂ emissions in eutrophic ecosystems.
1641 Besides, we evaluated the potential feedbacks between eutrophication and climate warming of
1642 a natural phytoplankton community and how increasing temperatures could affect the
1643 metabolic balance. For instance, we found that under no nutrient limitation, warming
1644 promoted an increase in productivity, with a dominance of cyanobacteria (with less overall
1645 diversity), and greater resource use efficiency. More interestingly, the ecosystem's metabolic
1646 balance changed, taking the microlakes in the direction of being CO₂ sources to the
1647 atmosphere. This finding gives experimental evidence of a positive feedback between
1648 eutrophication symptoms (cyanobacteria blooms) and warming, via higher CO₂ emission rates
1649 in cyanobacteria dominated warmer systems, adding to current research highlighting such
1650 self-reinforcing feedbacks.

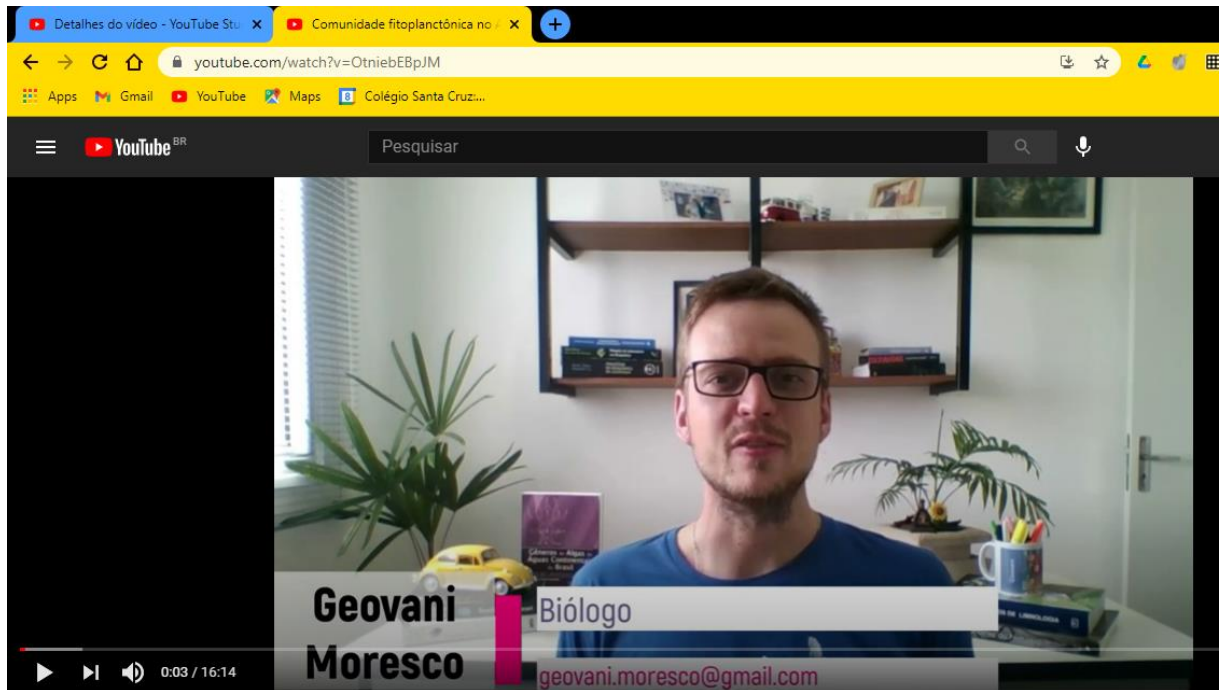
1651 In the second paper, we verified how cumulative stressors climate change (warming,
1652 eutrophication, and extremes rainfall events) can affect community and ecosystem stability.
1653 To do that, we used three phytoplankton communities already established and adapted to three
1654 different temperature scenarios. Our results showed experimentally that, environments
1655 stressed by warming combined with extreme rainfall events, seem to have a greater capacity
1656 to recover to the stage before the disturbance, such as restoring their biomass production and
1657 resource use efficiency. Thus, eutrophication symptoms are exacerbated by rising

1658 temperatures and are resilient to rainfall extreme events, reinforcing the dominance of
1659 cyanobacteria blooms.

1660 In conclusion, we experimentally demonstrated the effects of climate change
1661 components and the vulnerability of the freshwater phytoplankton community as well as its
1662 enormous potential to test relevant ecological hypotheses in relatively simple experiments.
1663 We used a highly diverse initial community and isolated it from potentially confounding
1664 factors. Thus, we were able to capture a pure pelagic response to climate change stressors,
1665 exclusive to the phytoplankton community, and how such responses translated to ecosystem
1666 functions (e.g., biomass production, energy transfer, and carbon cycle).

1667

1668 **APPENDIX A - Scientific Dissemination**
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1672 Fig S1. Scientific dissemination of the main results of the thesis for the participation of
1673 society on the theme of climate change. Video available at:
1674 <https://www.youtube.com/watch?v=OtniebEBpJM>

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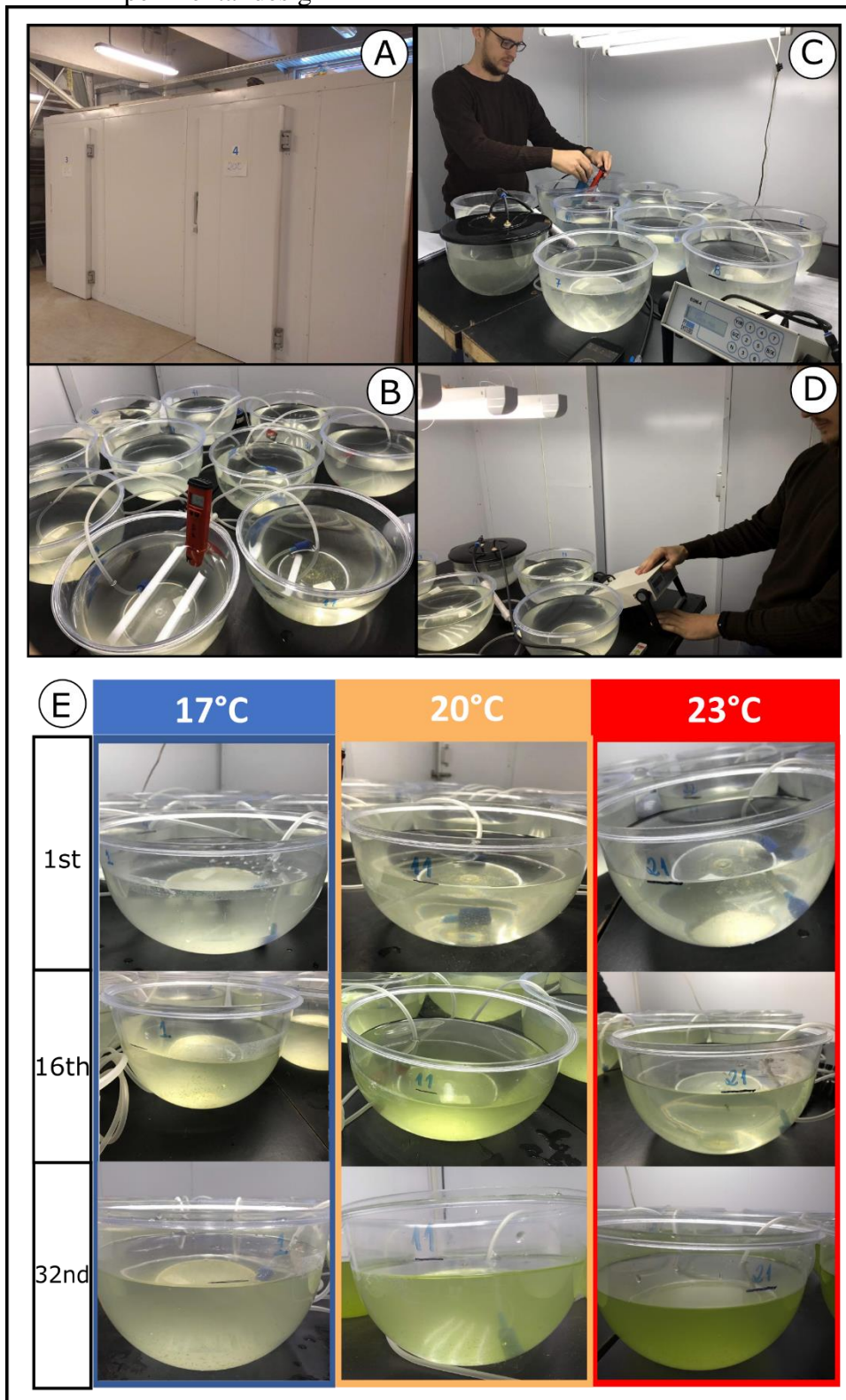
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1690 APPENDIX B - Experimental design

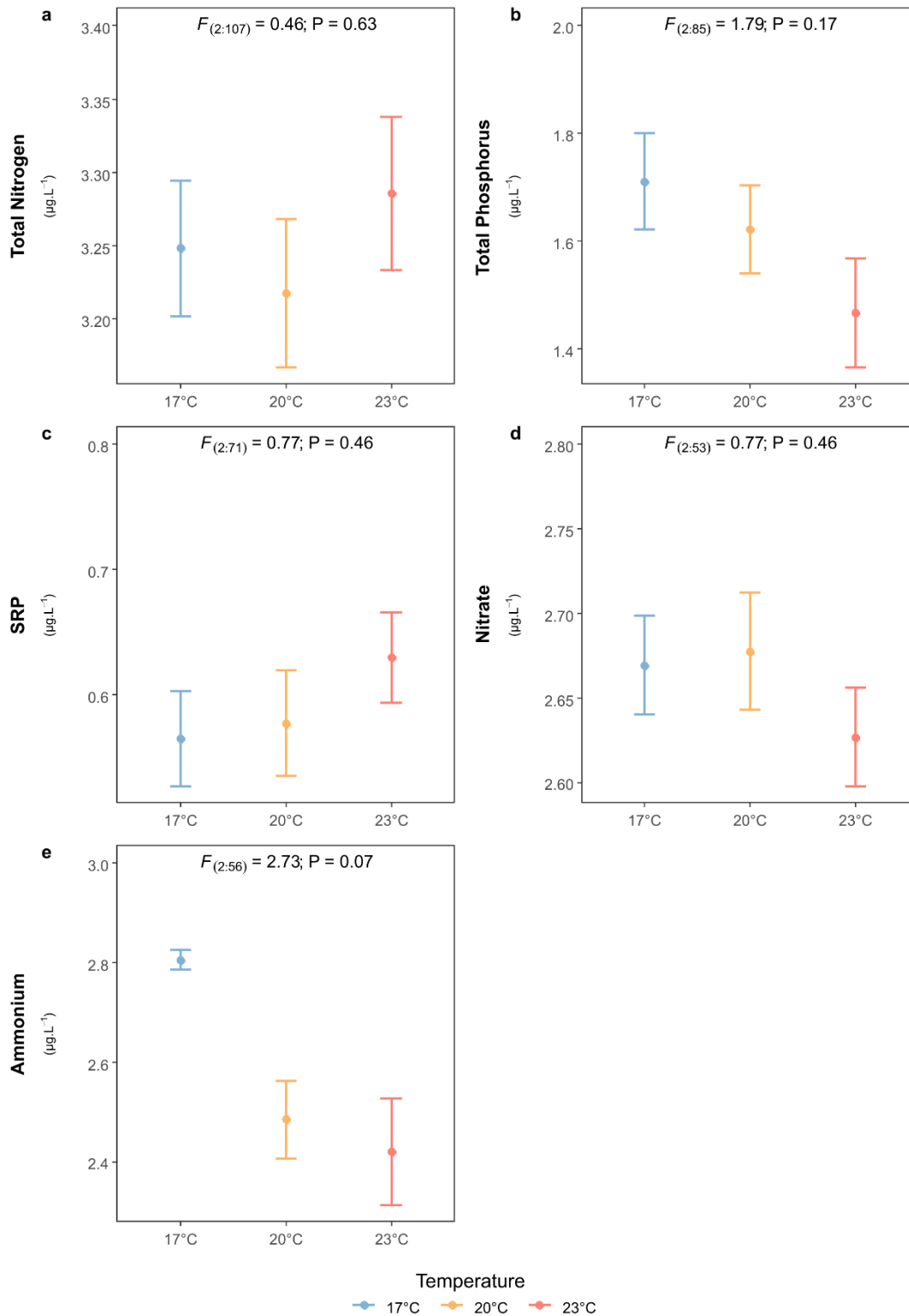


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1692 **Figure S1.** Experimental Design. A) Rooms temperature; B) overview of the experiment; C)
 1693 Monitoring of limnological variables; D) CO₂ fluxes measured; E) Development of the
 1694 phytoplankton community at the three temperatures evaluated (17°C, 20°C, and 23°C). Photos
 1695 were taken on the 1st day, on the 16th day (when temperatures reached the desirable), and on
 1696 the 32nd day.

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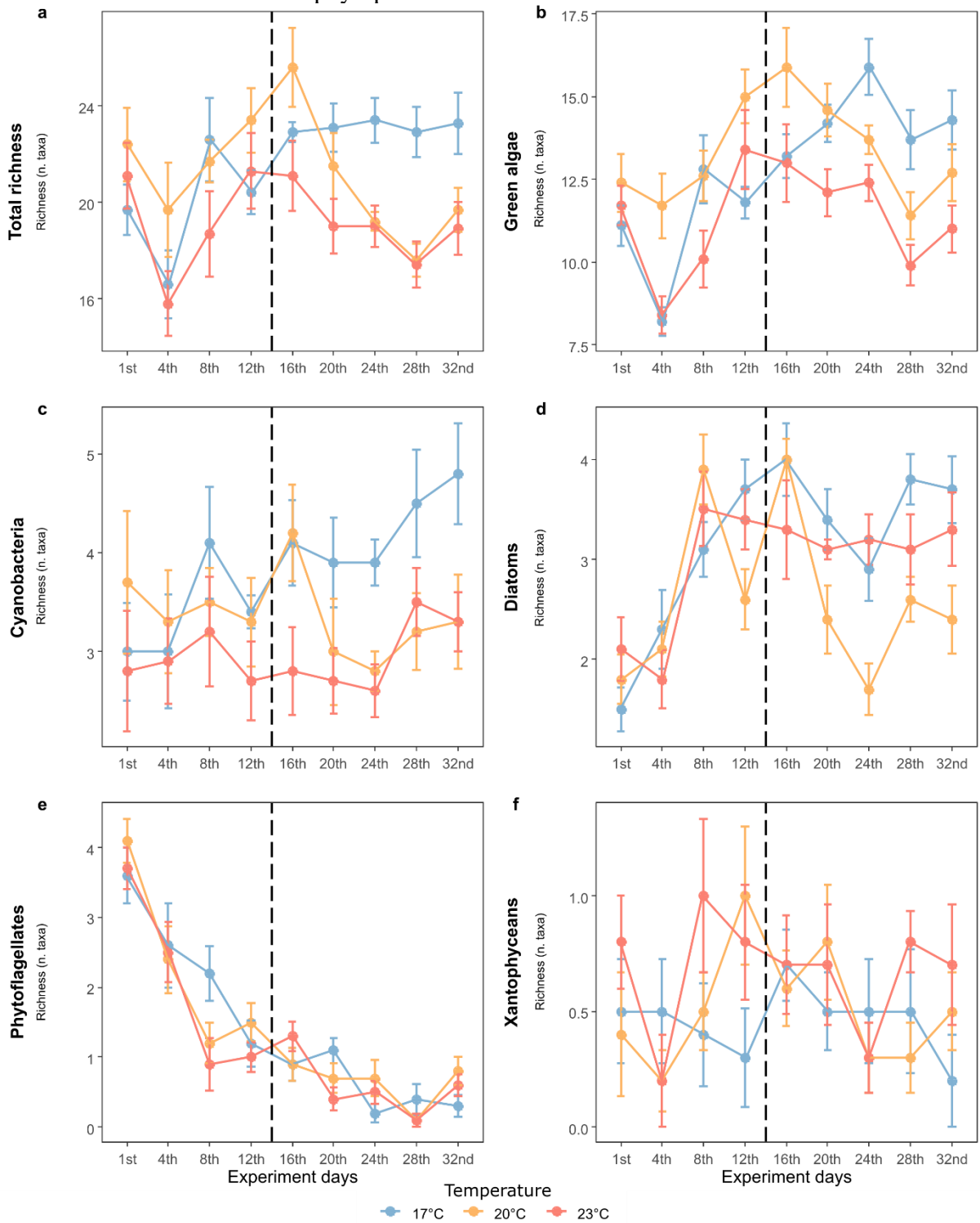
1698 APPENDIX C – Nutrient concentrations



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1700 **Figure S2:** Nutrient concentrations in the three temperature treatments obtained. The non-
 1701 significant result in the analysis of variance ($P > 0.05$) indicated that nutrient concentrations
 1702 were similar between treatments. Note that the scale varies among limnological variables. The
 1703 central point denotes the mean value and, whiskers represent standard error for each
 1704 temperature.

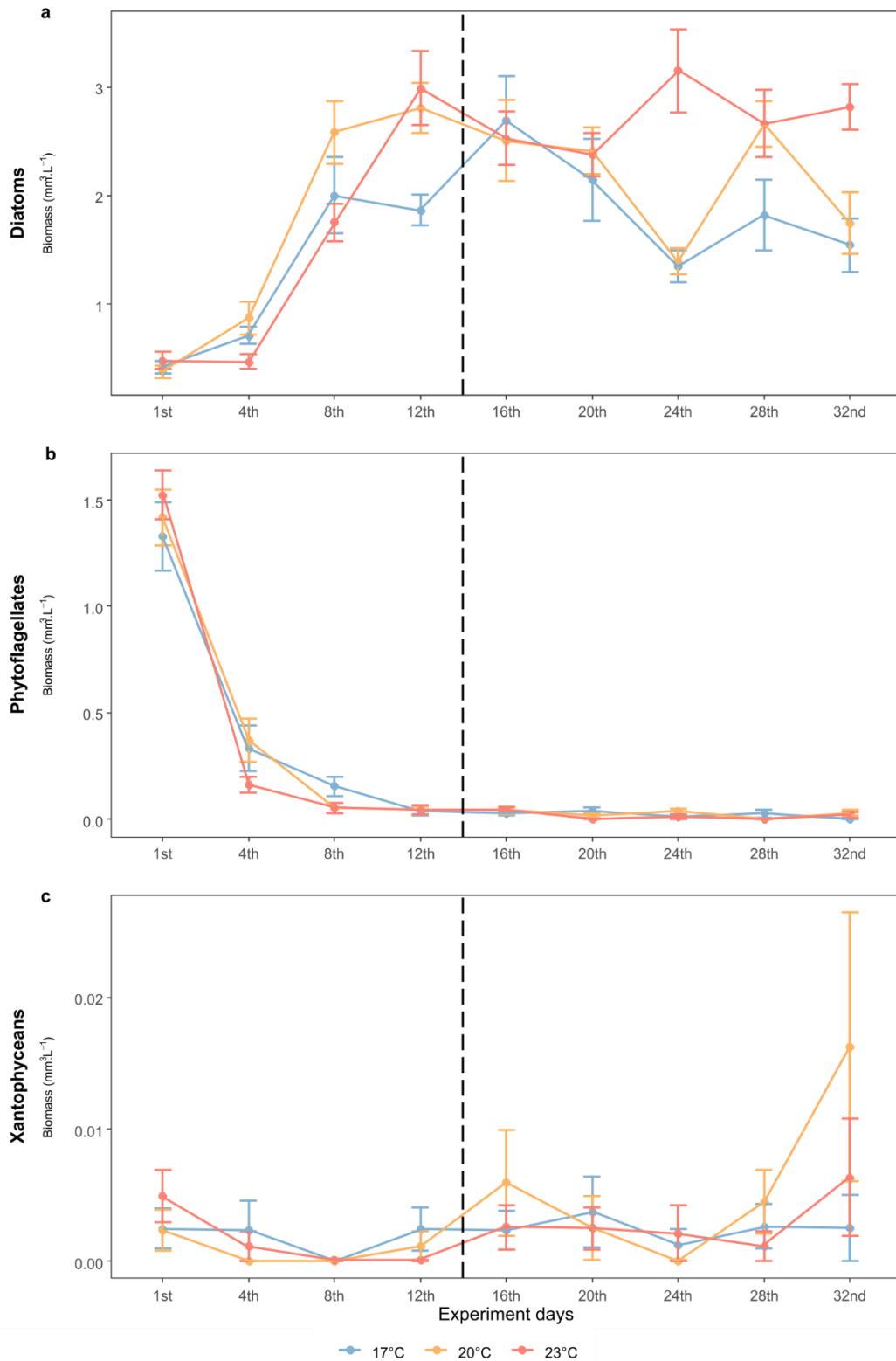
1705 APPENDIX D - Variation of phytoplankton taxonomic richness



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1707 **Figure S3:** Variation of phytoplankton taxonomic richness with temperature through the
 1708 32 days of the experiment: total (a), green algae (b), cyanobacteria (c), diatoms (d),
 1709 phytoplagellates (e) and xanthophyceans (f). The dotted line indicates that on the 16th
 1710 day of the experiment the wished temperatures were reached. The central point denotes
 1711 the mean value and, whiskers represent standard error for each temperature in each
 1712 experiment day.
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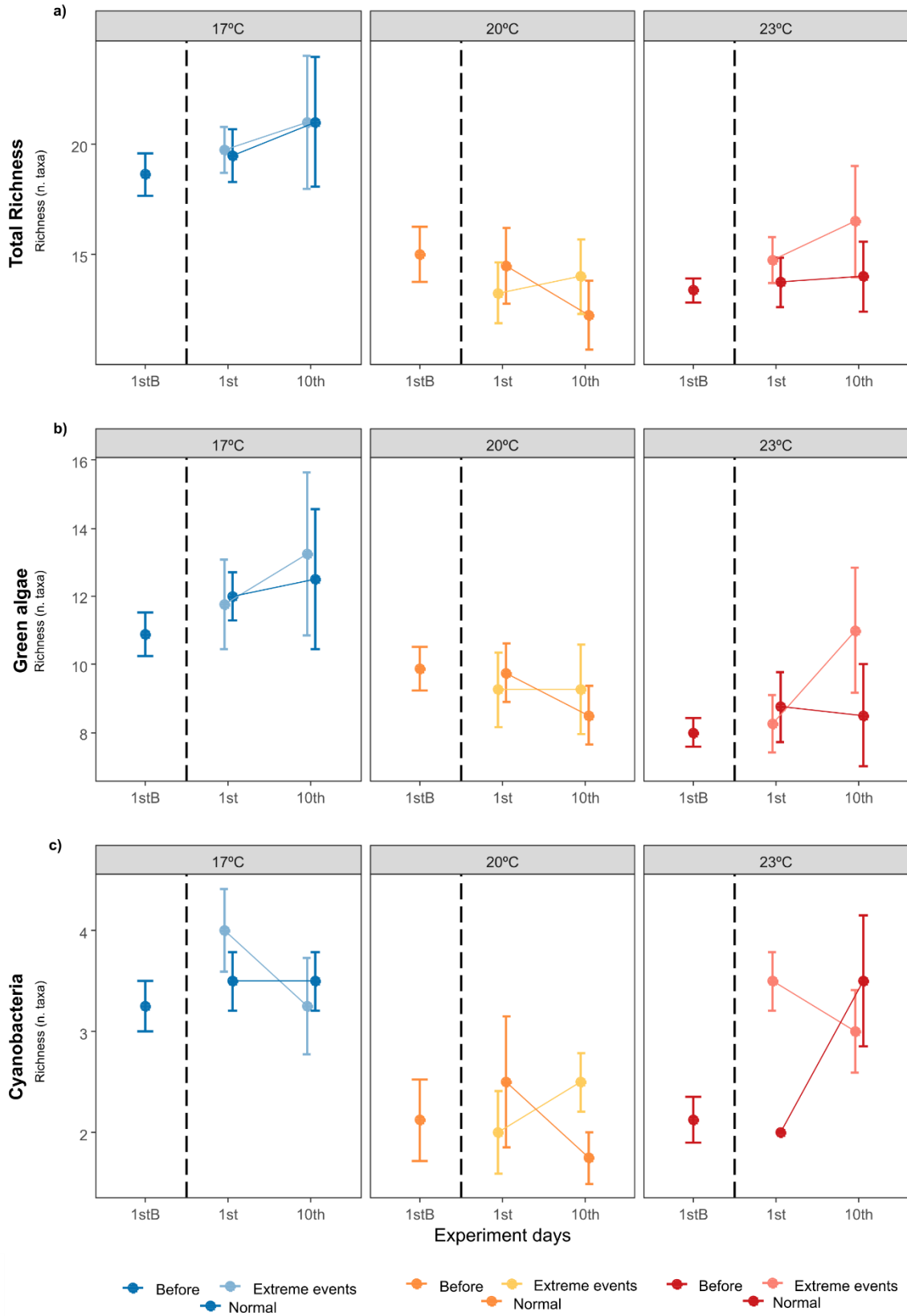
1714 APPENDIX E - Variation of phytoplankton biomass
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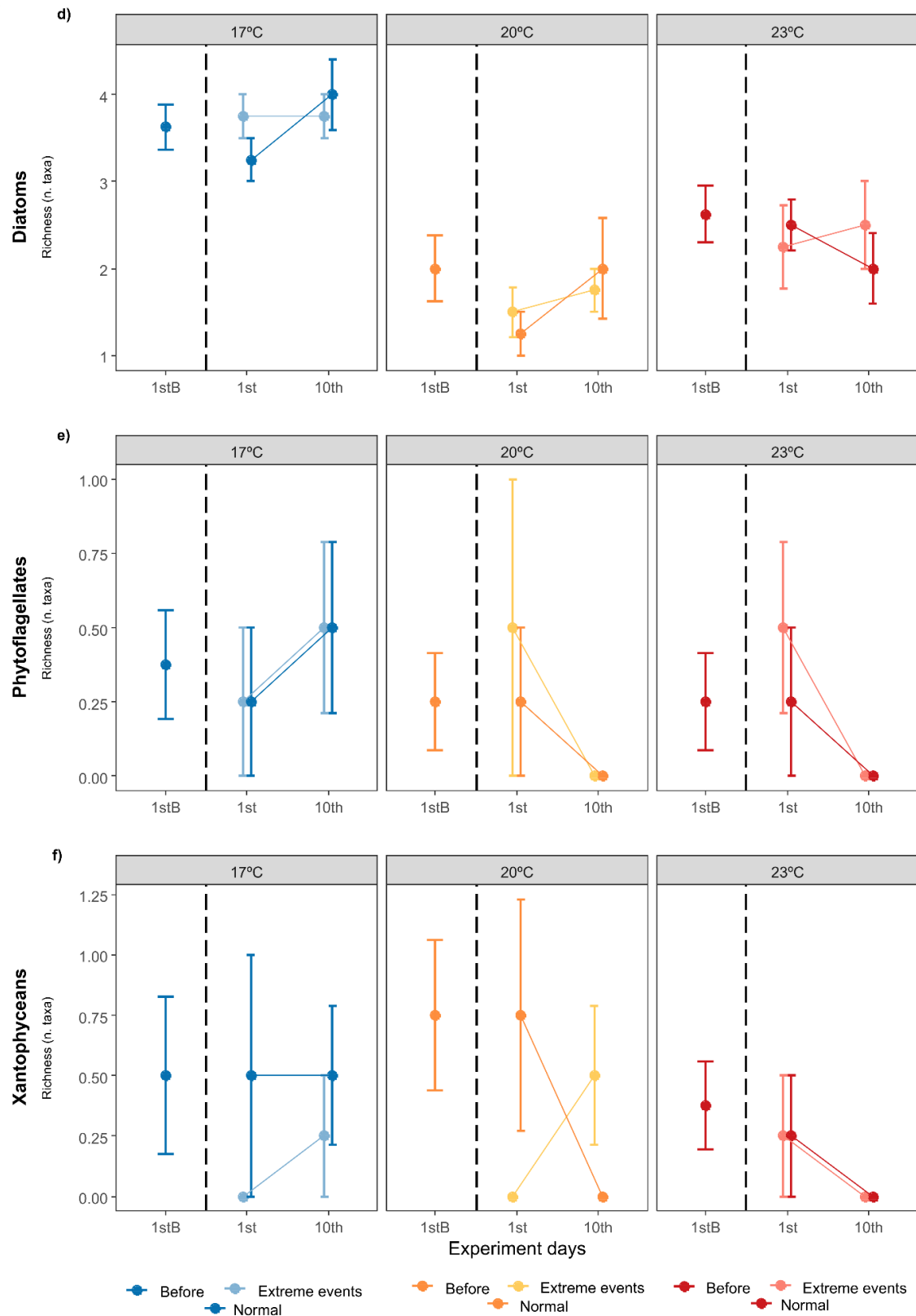
1717 **Figure S4:** Variation of phytoplankton biomass with temperature through the 32 days of the
1718 experiment: diatoms (a), phytoflagellates (b) and xantophyceans (c). The dotted line indicates
1719 that on the 16th day of the experiment the wished temperatures were reached. The central point
1720 denotes the mean value and, whiskers represent standard error for each temperature in each
1721 experiment day.

1722 APPENDIX F - Variation of phytoplankton taxonomic richness
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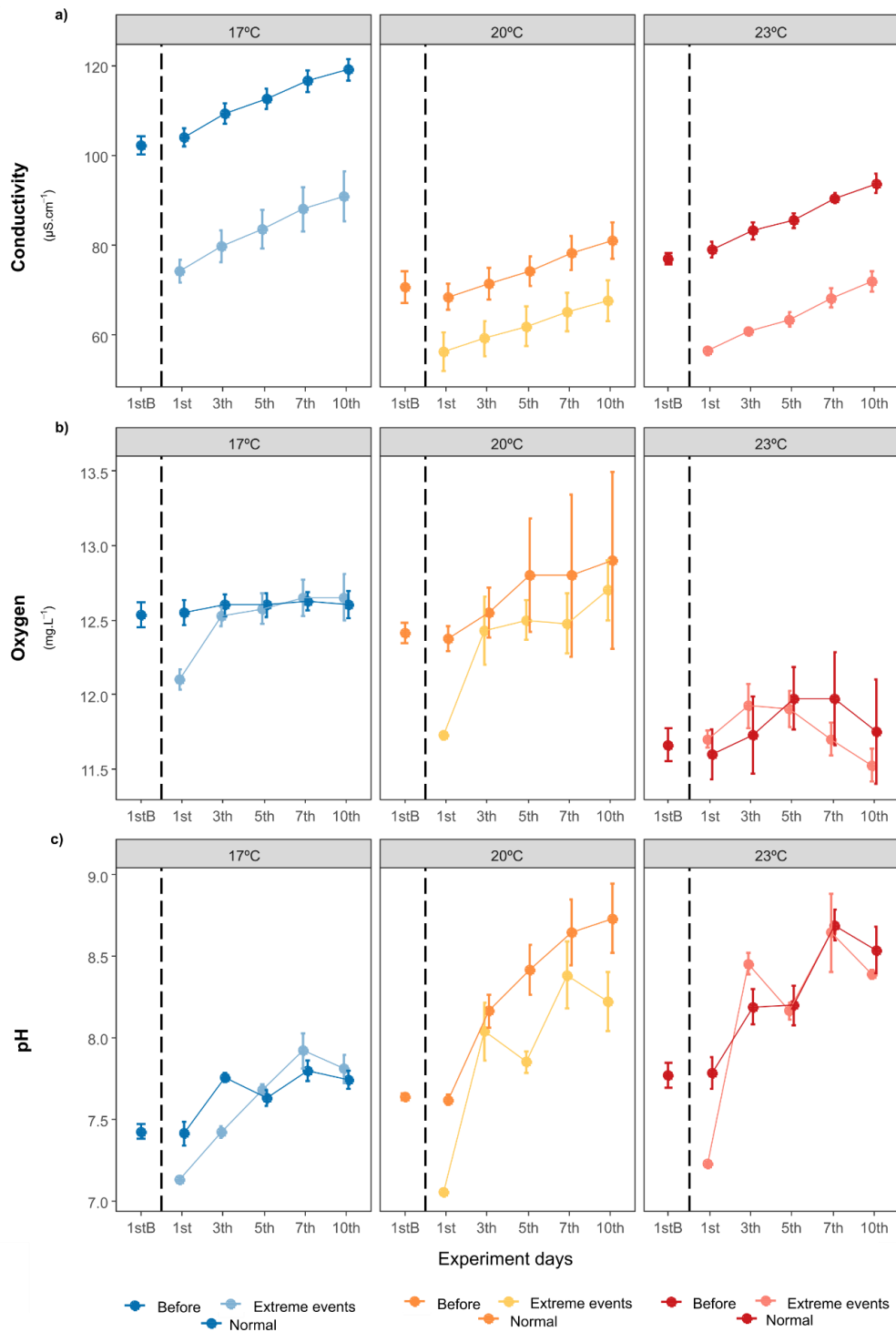
1726

1727 **Supplementary Figure 1:** Variation of phytoplankton taxonomic richness with temperature through
 1728 the 10 days of the experiment: total (a), green algae (b), cyanobacteria (c), diatoms (d),
 1729 phytoflagellates (e), xanthophyceans (f). The dotted line indicates the occurrence of the extreme
 1730 rainfall event (1stB). Light colors indicate communities subject to extreme rainfall and dark colors the

1731 undisturbed communities. The central point denotes the mean value and, whiskers represent standard
 1732 error.

1733 **APPENDIX G - Variation of abiotic variables**

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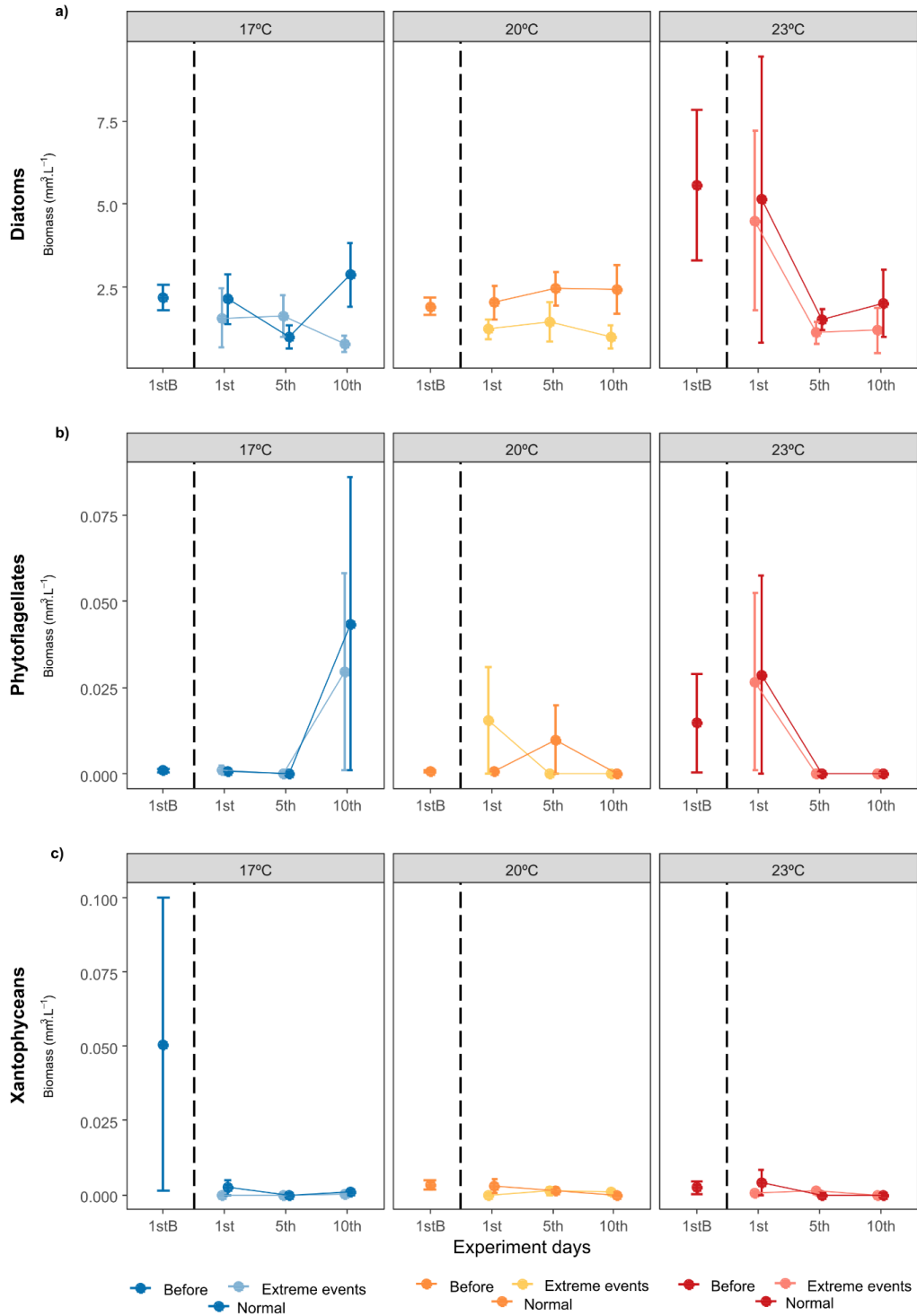


1735

1736 **Supplementary Figure 2:** Variation of abiotic variables in the three temperature treatments (17
 1737 °C, 20 °C, and 23 °C) through the 10 days of the experiment. The dotted line indicates the
 1738 occurrence of the extreme rainfall event (1stB). Light colors indicate communities subject to

1739 extreme rainfall and dark colors the undisturbed communities. The central point denotes the
 1740 mean value and, whiskers represent standard error.

1741 **APPENDIX H - Variation of phytoplankton biomass**
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1744 **Supplementary Figure 3.** Variation of phytoplankton biomass with temperature
 1745 through the 10 days of the experiment: diatoms (a), phytoflagellates (b),
 1746 xantophyceans (c). The dotted line indicates the environments before the period of
 1747 extreme rainfall event (1stB) of the period after the disturbance. Light colors indicate

- 1748 environments with the effect of extreme rainfall and dark colors without the effect. The
1749 central point denotes the mean value and, whiskers represent standard error.
- 1750 **ANNEX A** - Research papers accepted or published during the doctoral development period
1751 that contributed to the execution of this thesis
1752
- 1753 Dunck, B., M. G. Junqueira, M. V. da Silva, and others. 2018. Periphytic and planktonic algae records
1754 Dunck, B., M. G. Junqueira, M. V. da Silva, A. Pineda, A. C. M. de Paula, B. F. Zanco,
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